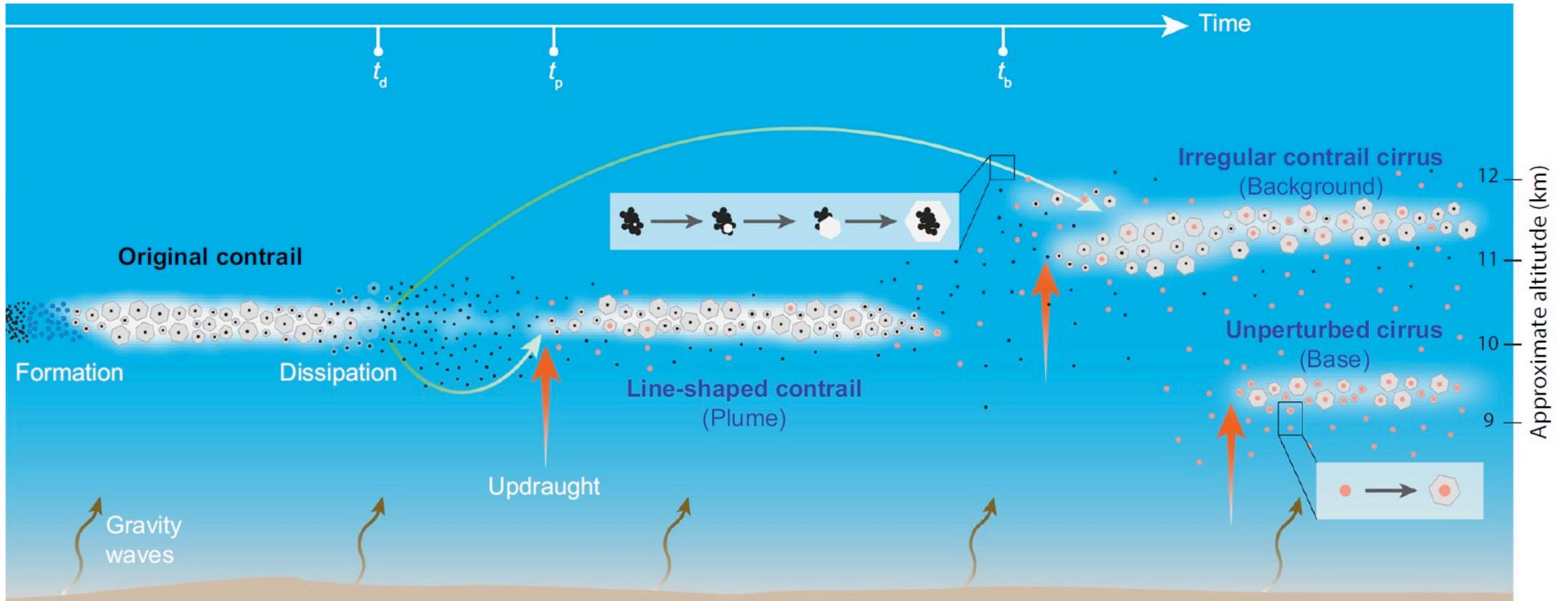


# Heterogeneous ice nucleation of soot particles: measurements, predictions and implications

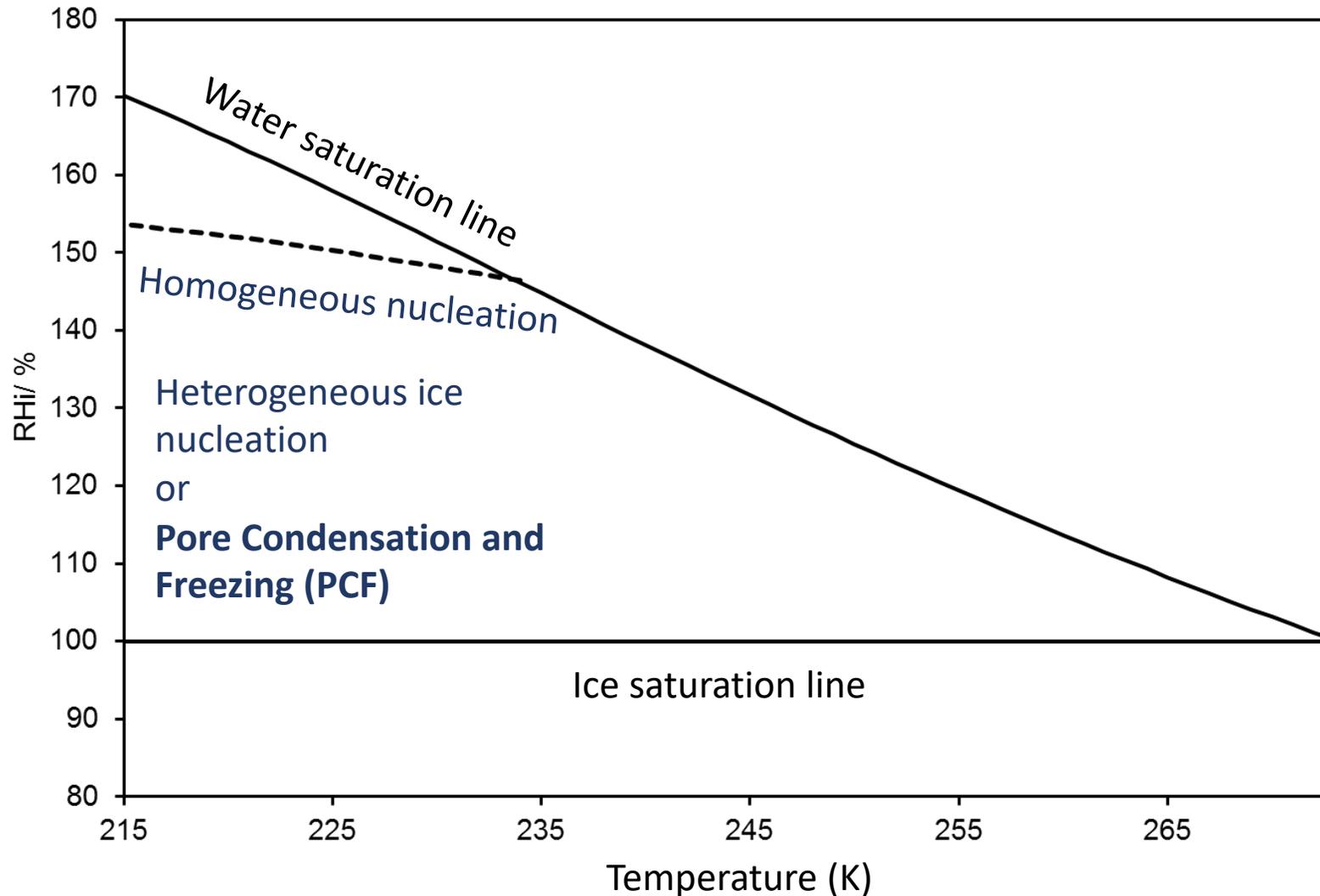
Claudia Marcolli<sup>1</sup>, Fabian Mahrt<sup>2</sup>, Bernd Kärcher<sup>3</sup>

<sup>1</sup>ETH Zürich, <sup>2</sup>Paul Scherrer Institute, <sup>3</sup>DLR Oberpfaffenhofen

# Overview



# RH/T conditions for the different ice nucleation modes



## **Homogeneous ice nucleation:**

Occurring on ubiquitous solution droplets

## **Heterogeneous ice nucleation:**

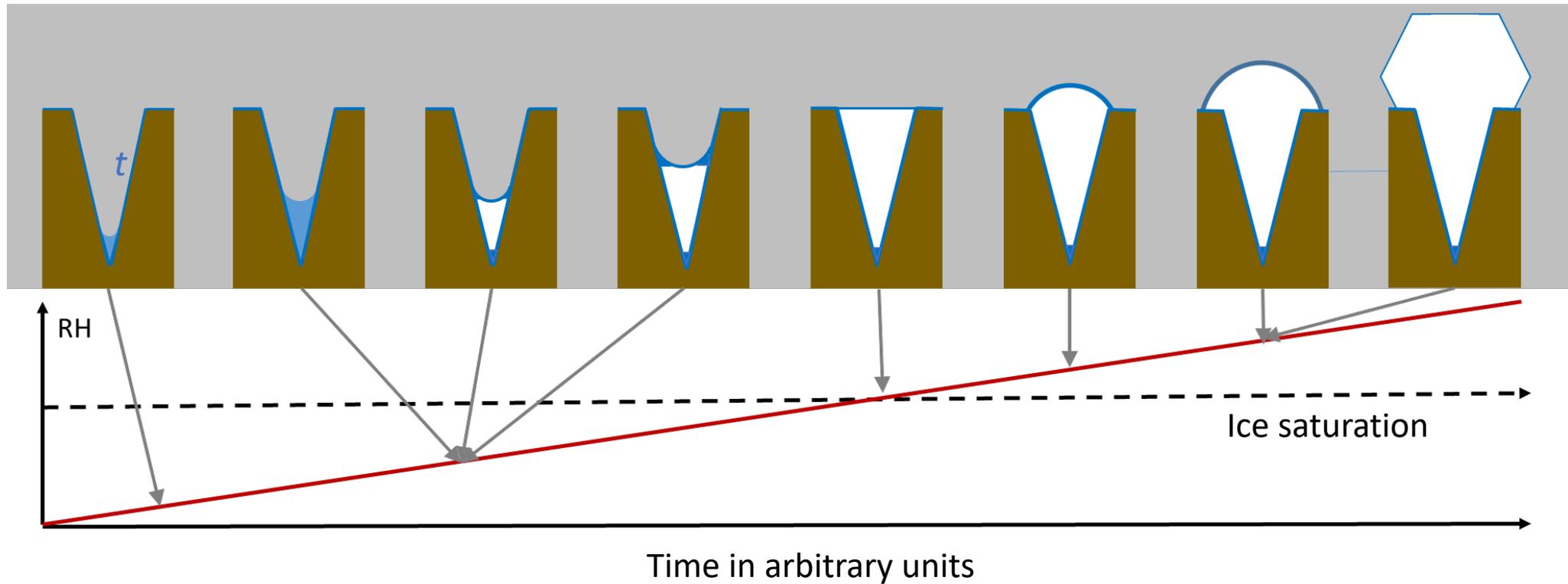
Occurring on ice-nucleating particles (INPs) with nucleation sites

## **PCF:**

Occurring on porous particles through homogeneous ice nucleation of pore water

# Pore condensation and freezing (PCF)

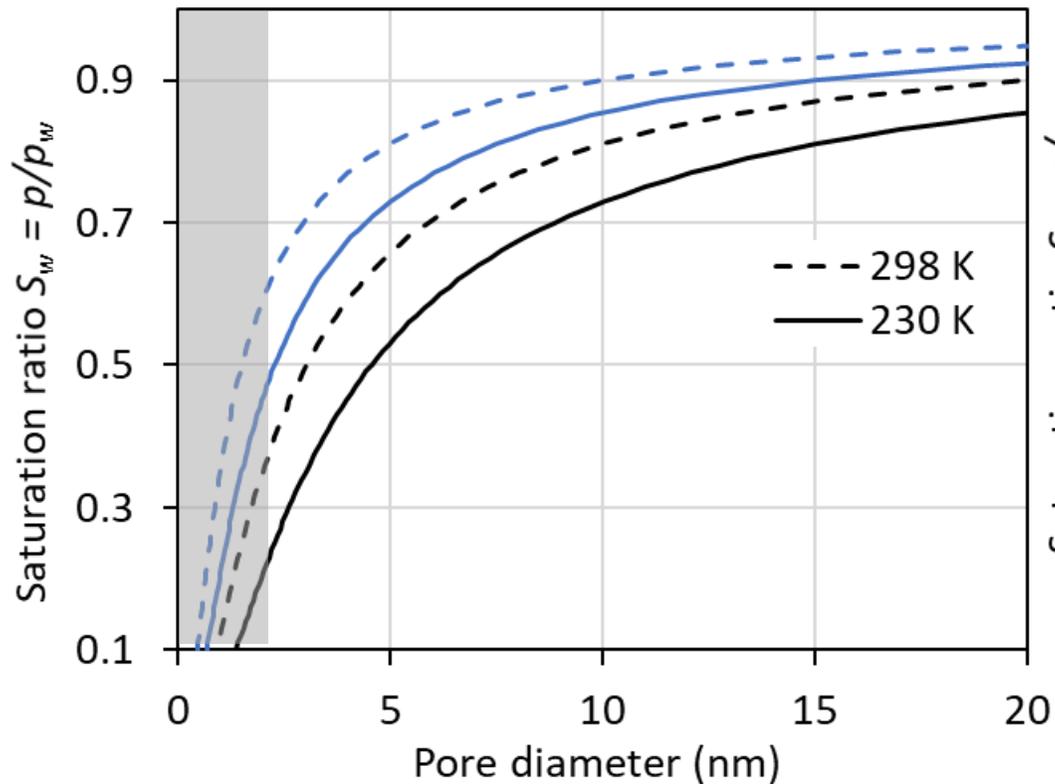
- Capillary condensation of water into pores (Kelvin equation)
- Nucleation of ice within the pores (Classical nucleation theory)
- Growth of ice out of the pores (Kelvin equation)



# What is the relevant diameter range for PCF?

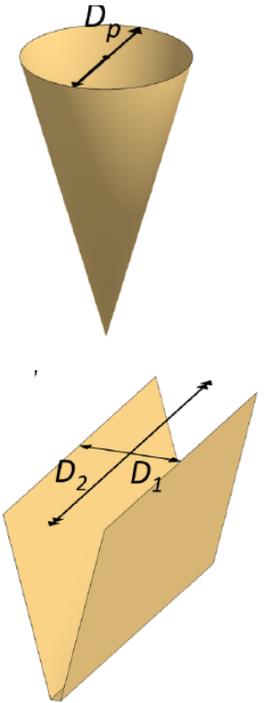
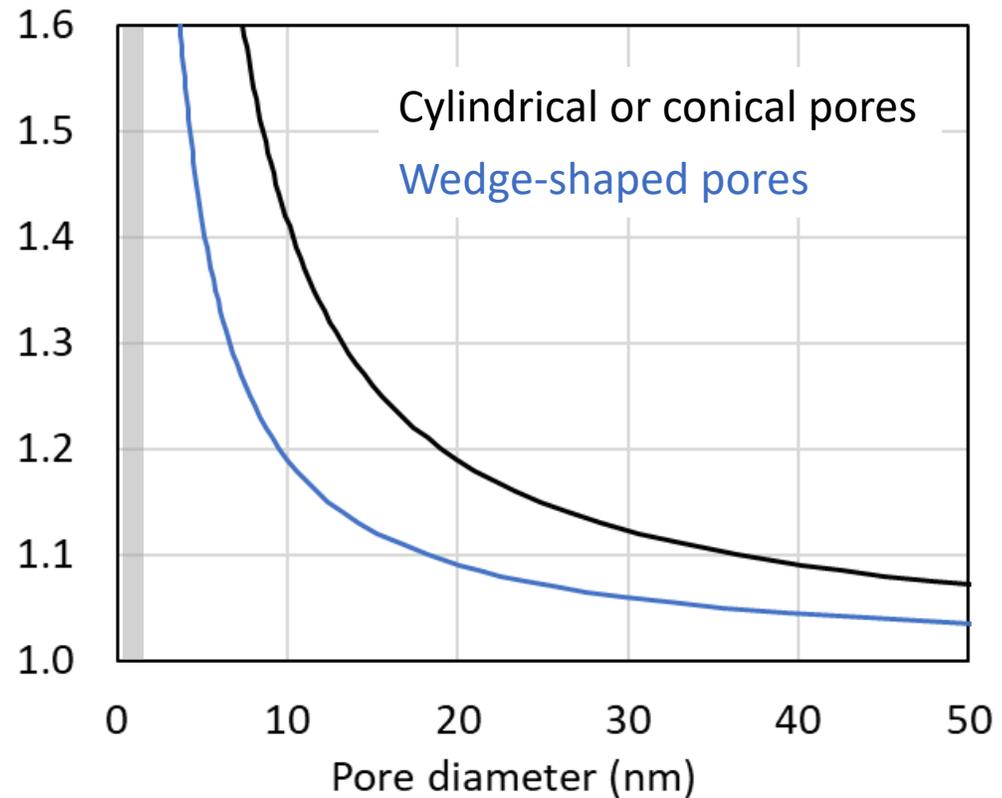
## Pore filling with water

Too narrow to freeze

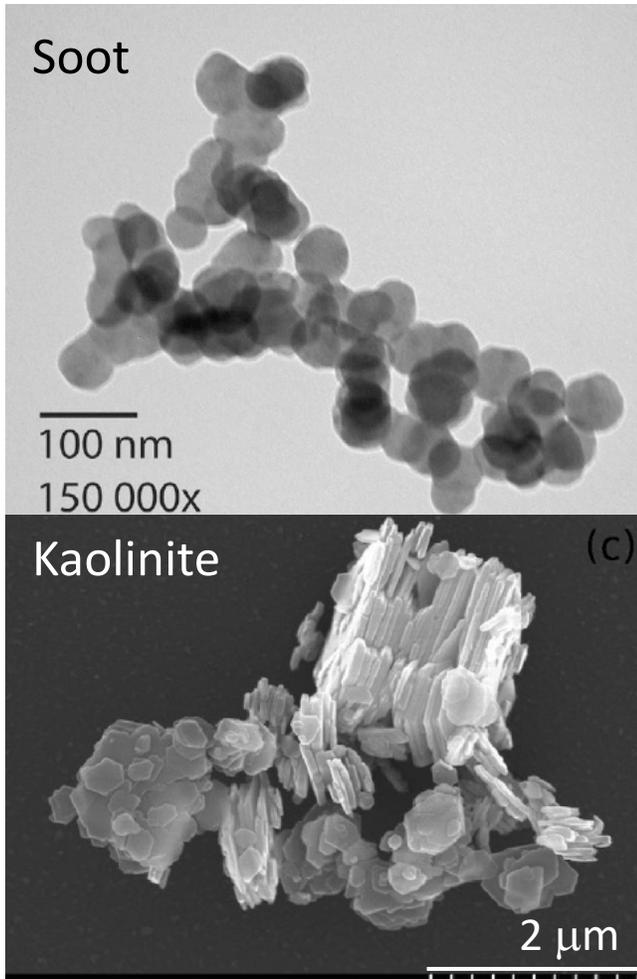


## Ice growth out of the pores

Too narrow to freeze



# What is the porosity of aerosol particles?



## **Soot**

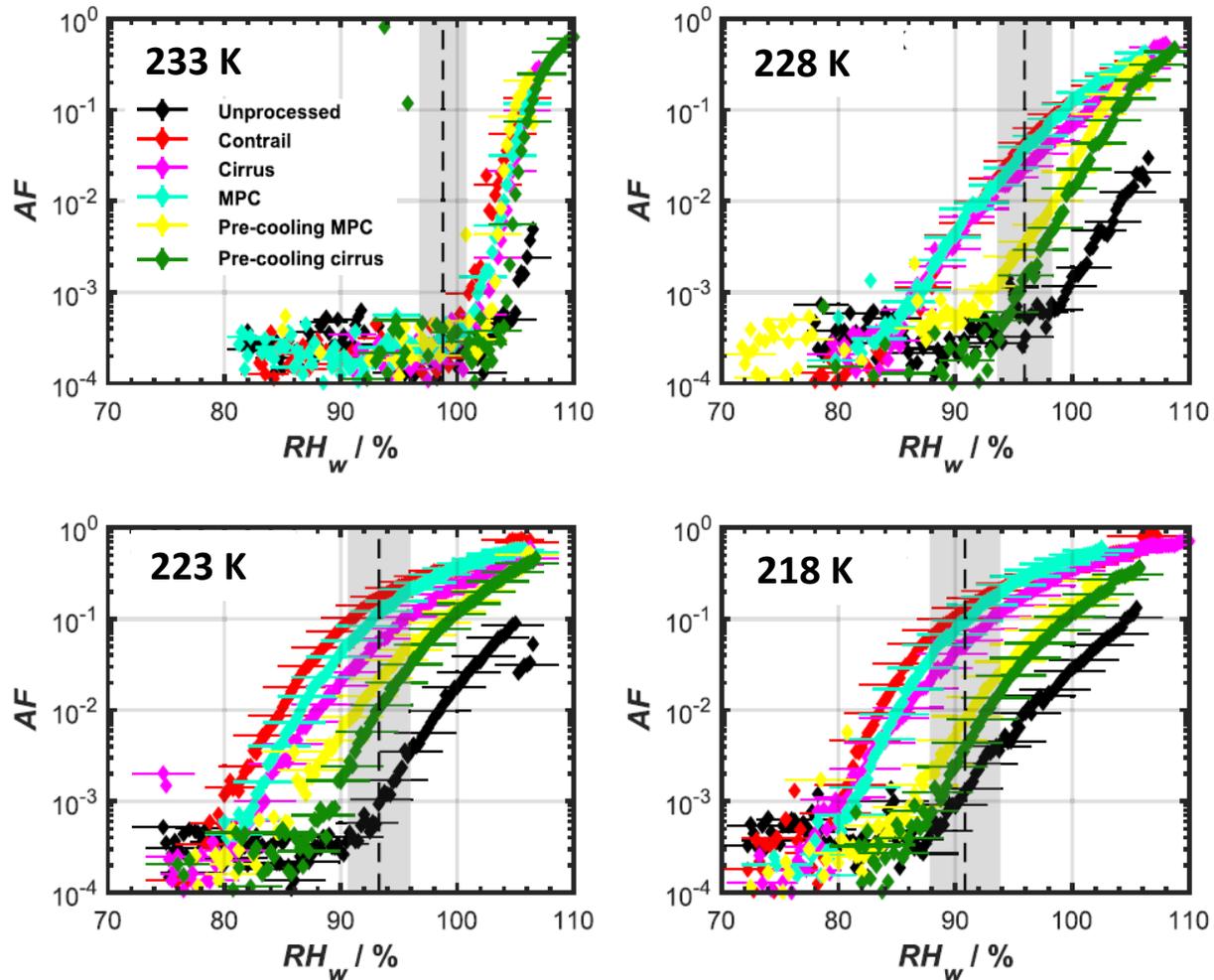
Porosity arising between primary particles

## **Mineral dust**

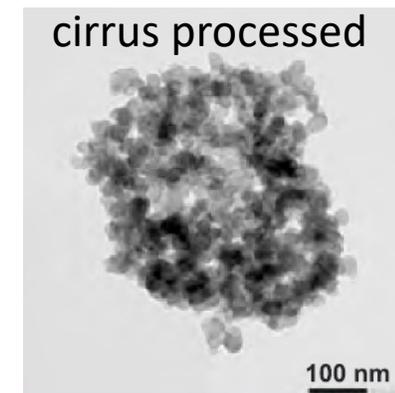
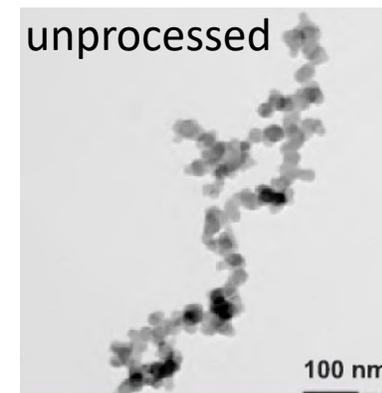
Porosity arising:

- in aggregates of crystallites
- At edges of clay minerals

# Activated fraction of processed miniCAST black

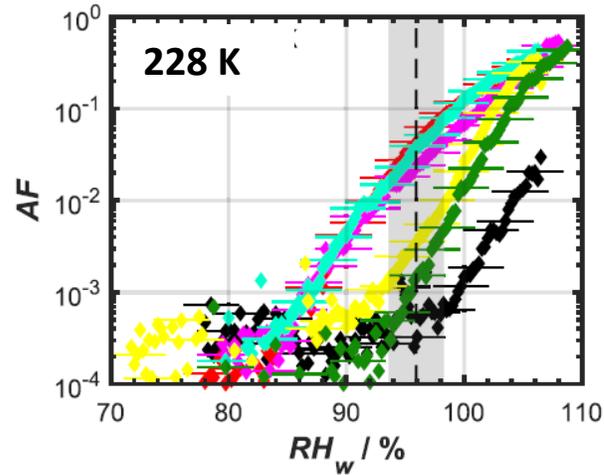
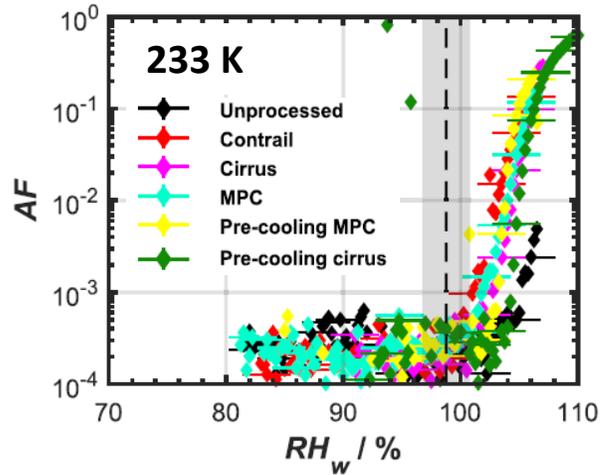


	Preprocessing	Evaporation
Unprocessed	—	—
contrail	104 % $RH_w$ 228 K	38 % $RH_w$ 233 K
cirrus	96 % $RH_w$ 228 K	38 % $RH_w$ 233 K
MPC	108 % $RH_w$ 243 K	30 % $RH_w$ 253 K
Precool MPC	96 % $RH_w$ 243 K	30 % $RH_w$ 253 K
Precool cirrus	65 % $RH_w$ 228 K	38 % $RH_w$ 233 K

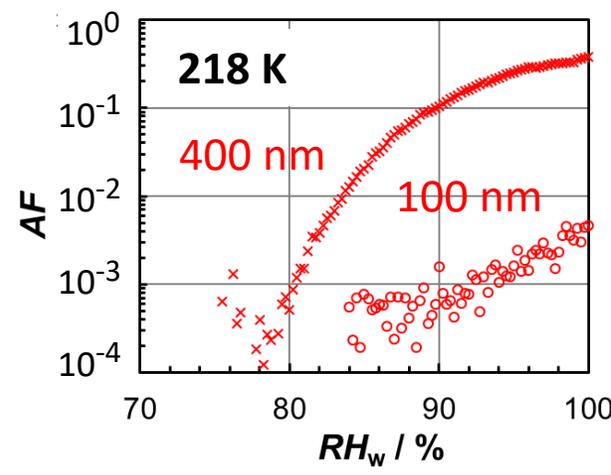
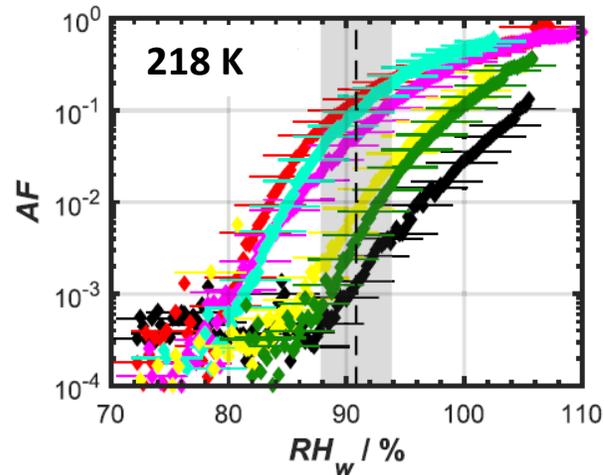
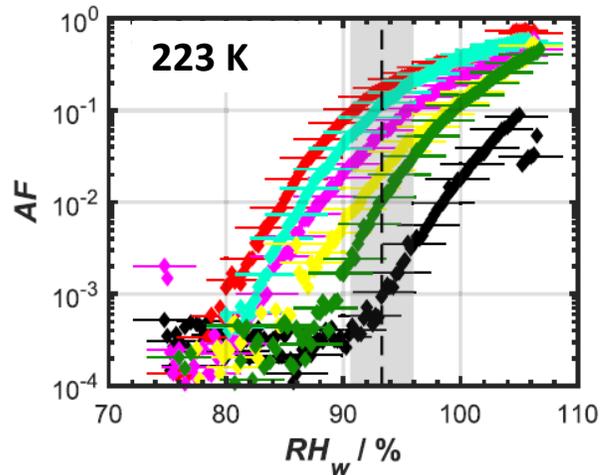
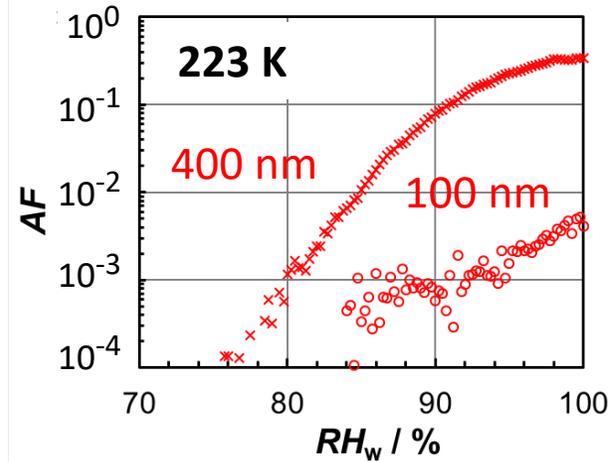


# Activated fraction of processed miniCAST black

400 nm diameter

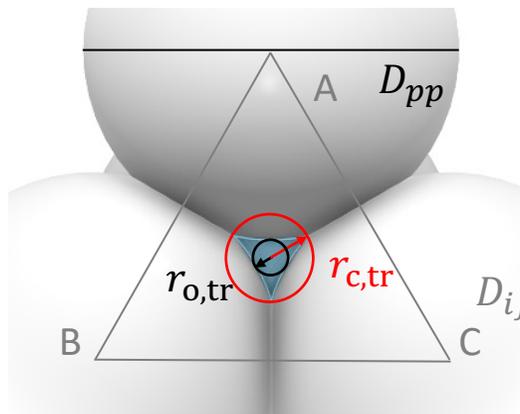
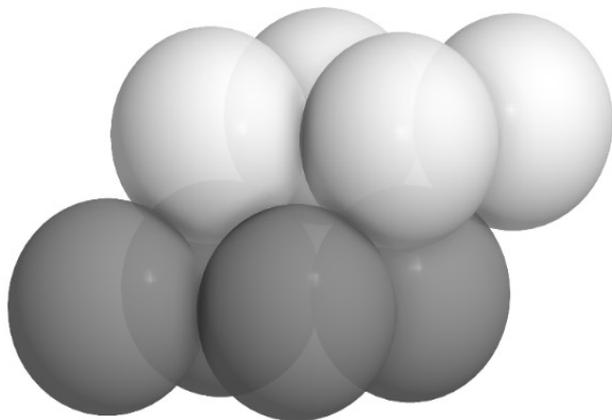


Size dependence



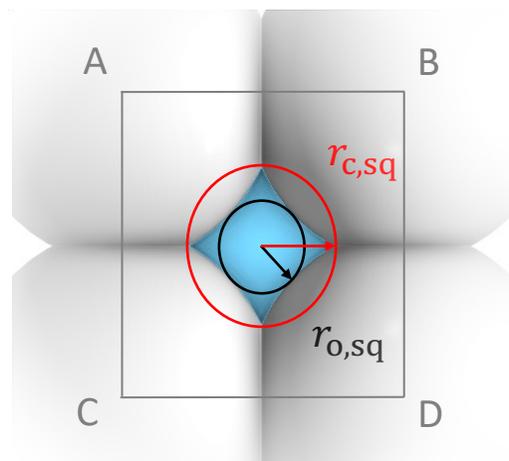
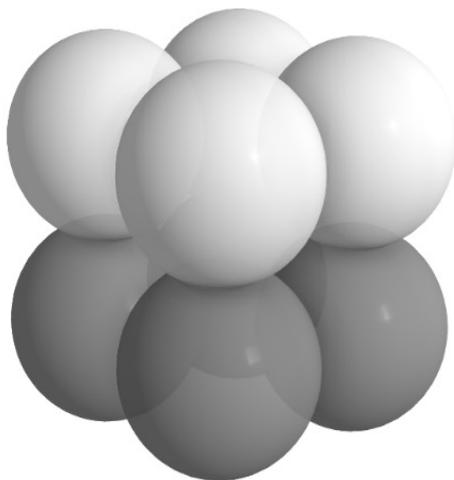
# Pores within soot aggregates

Tetrahedral packing



Three-membered ring pore

Cubic packing



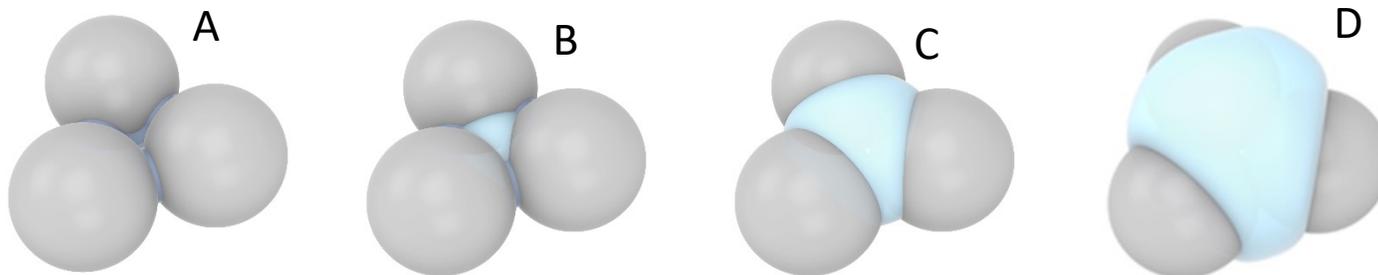
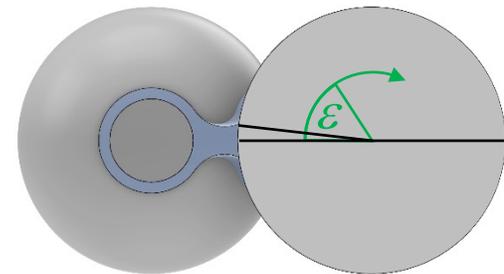
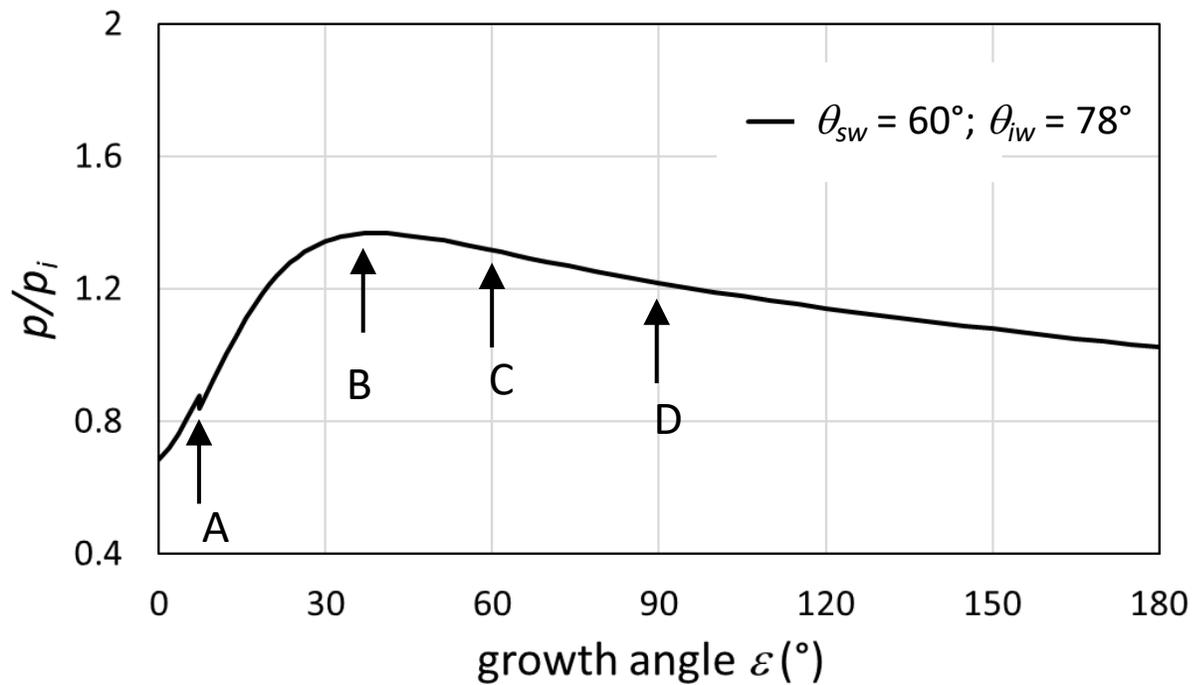
Four-membered ring pore

Spherical primary particles with typical diameters  $D_{pp} = 10\text{--}30$  nm

$$\text{Overlap } C_{ov} = \frac{D_{pp} - D_{ij}}{D_{pp}} = 0.01\text{--}0.2$$

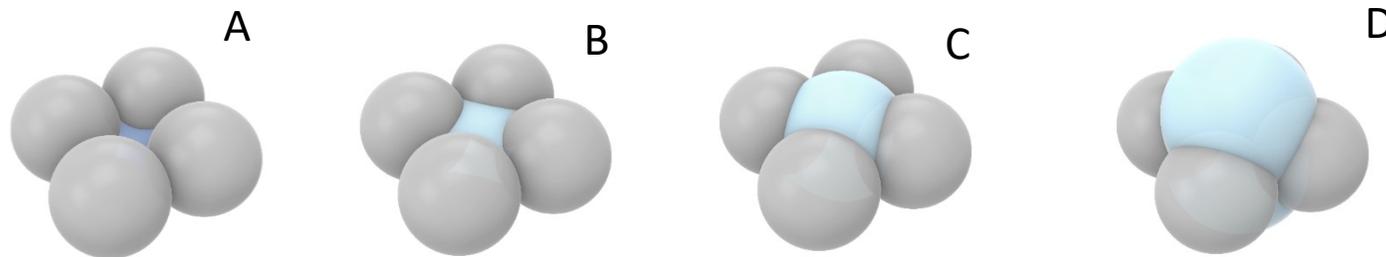
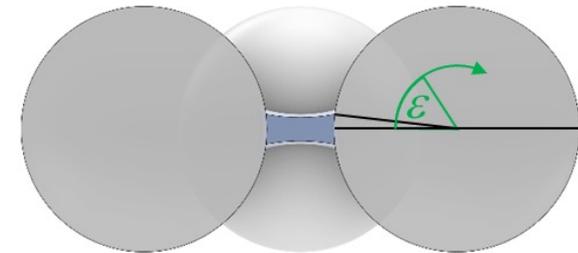
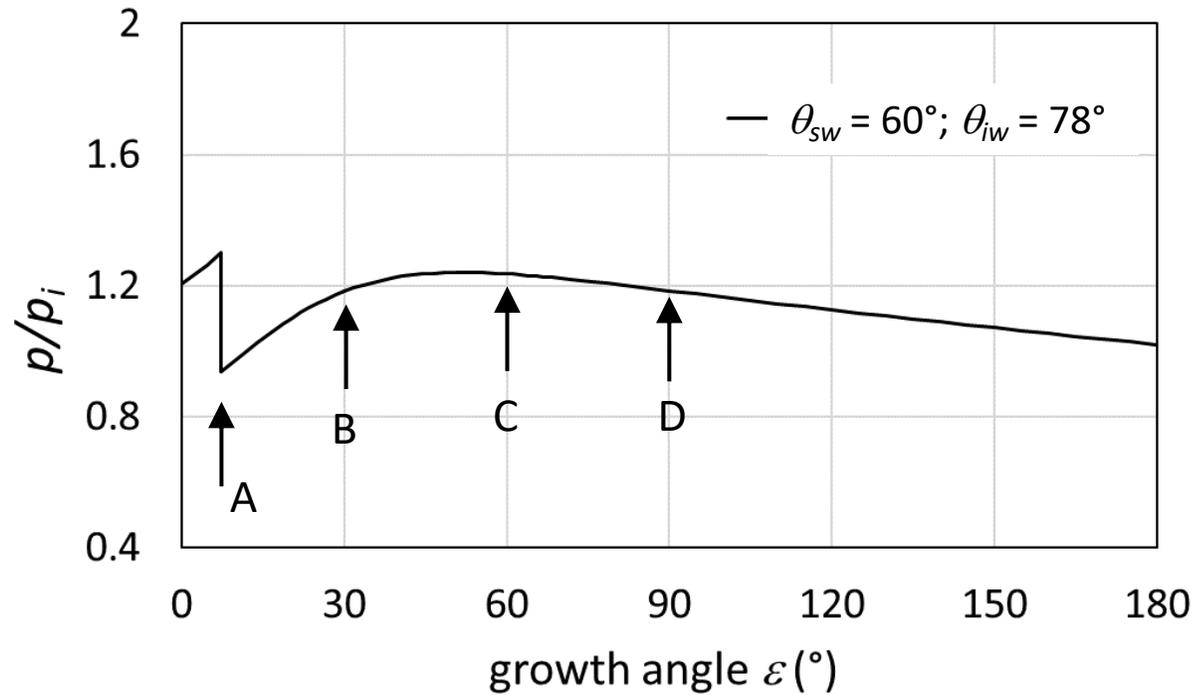
# Soot PCF on three-membered ring pore

primary particle diameter: 20 nm,  $C_{ov} = 0.05$

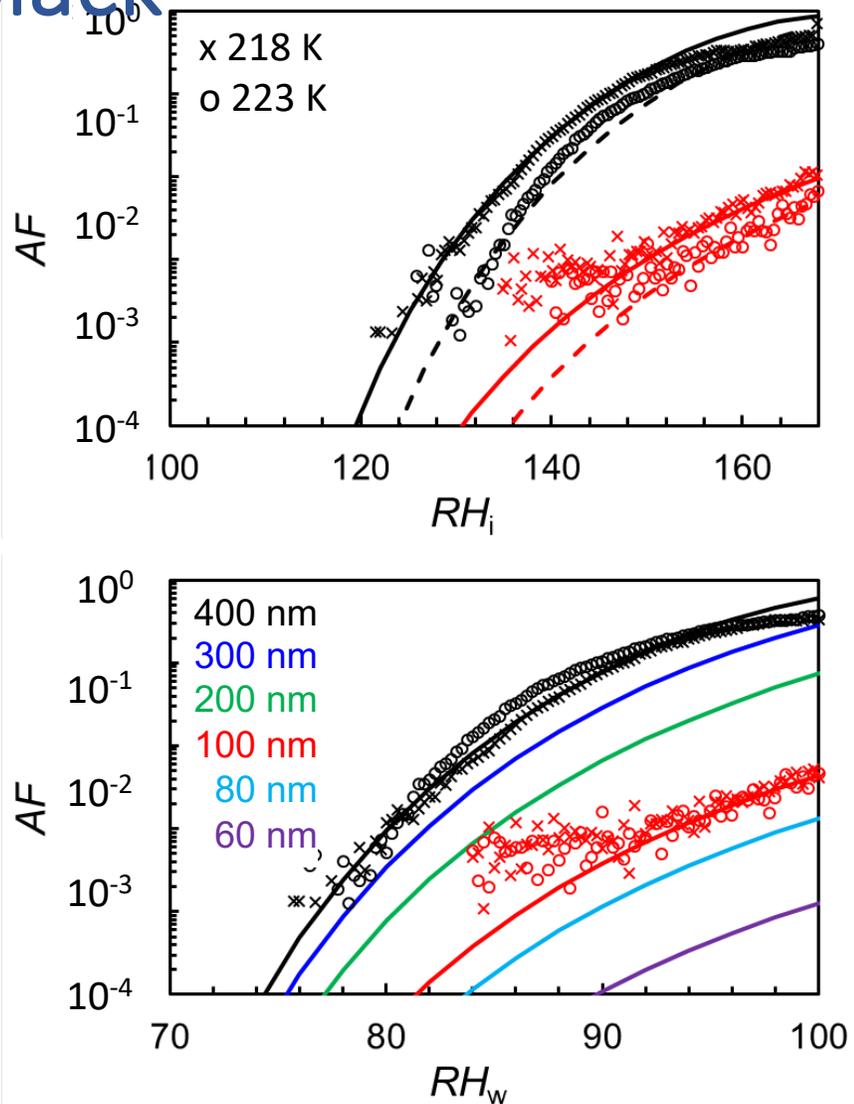


# Soot PCF on four-membered ring pore

primary particle diameter: 20 nm,  $C_{ov} = 0.1$



# Activated fraction (AF) parameterization of miniCAST black



$$AF = 1 - (1 - P_N(RH)) \left( (N_p - 2)^{1.86} \right)$$

$$P_N(RH_w) = 10^{\left( \frac{1}{0.3374 - 0.006091 RH_w} \right)}$$

Primary particle diameter:  $D_{pp} = 31 \text{ nm}$

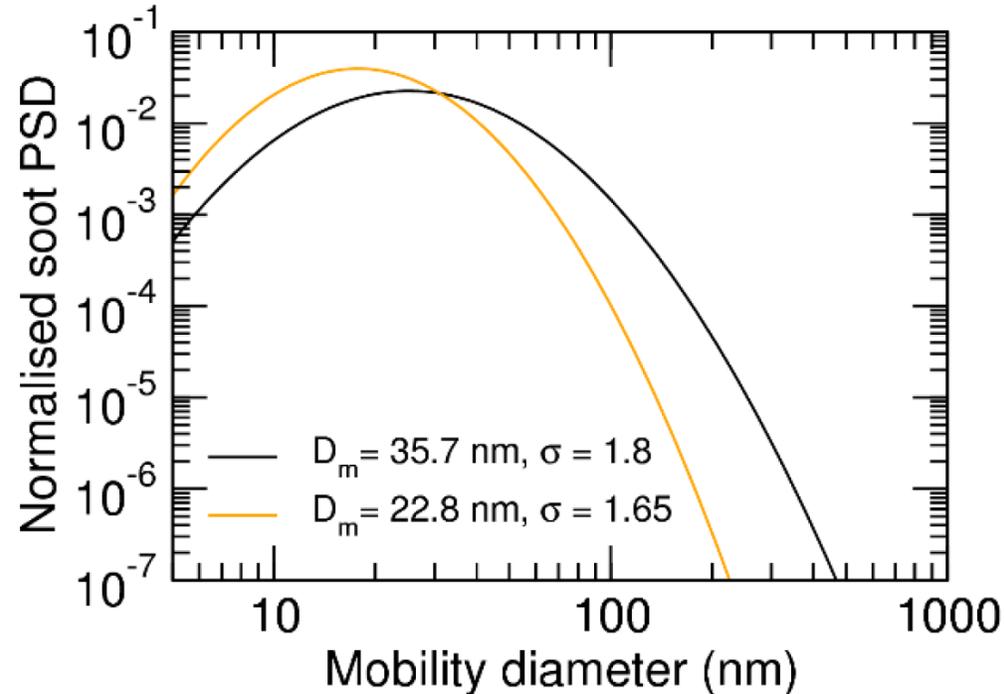
Fractal dimension: 1.86

→ Number of particles ( $N_p$ ) in aggregate with diameter  $D_m$

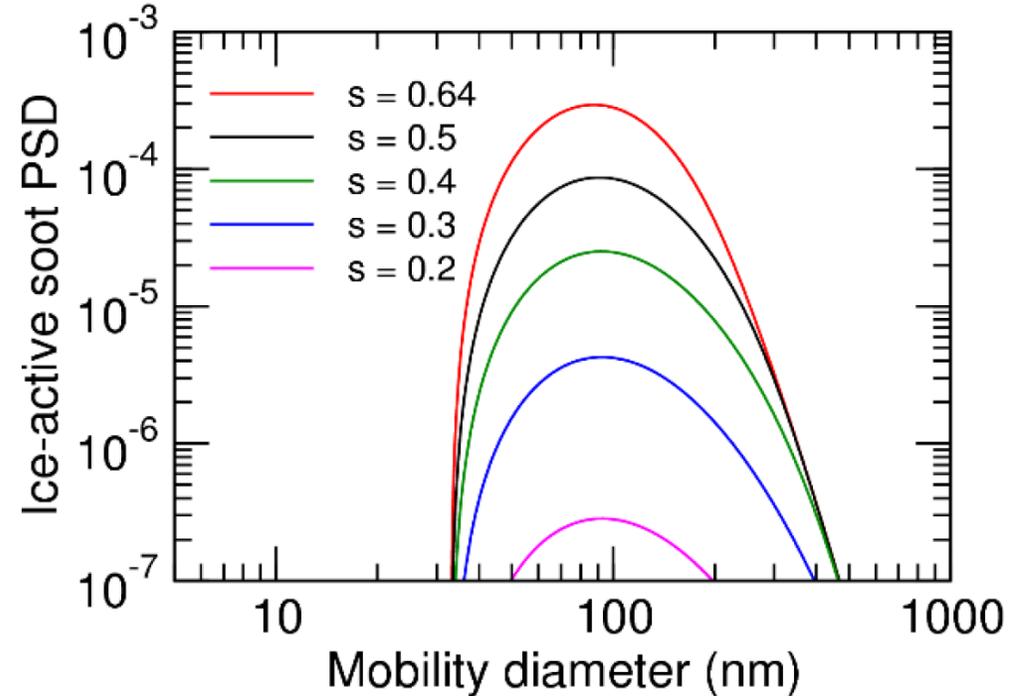
$D_m$	$N_p$
400 nm	94
300 nm	55
200 nm	26
100 nm	7
80 nm	5
60 nm	3

# PCF parameterization for aircraft soot

Aircraft soot particle size distribution



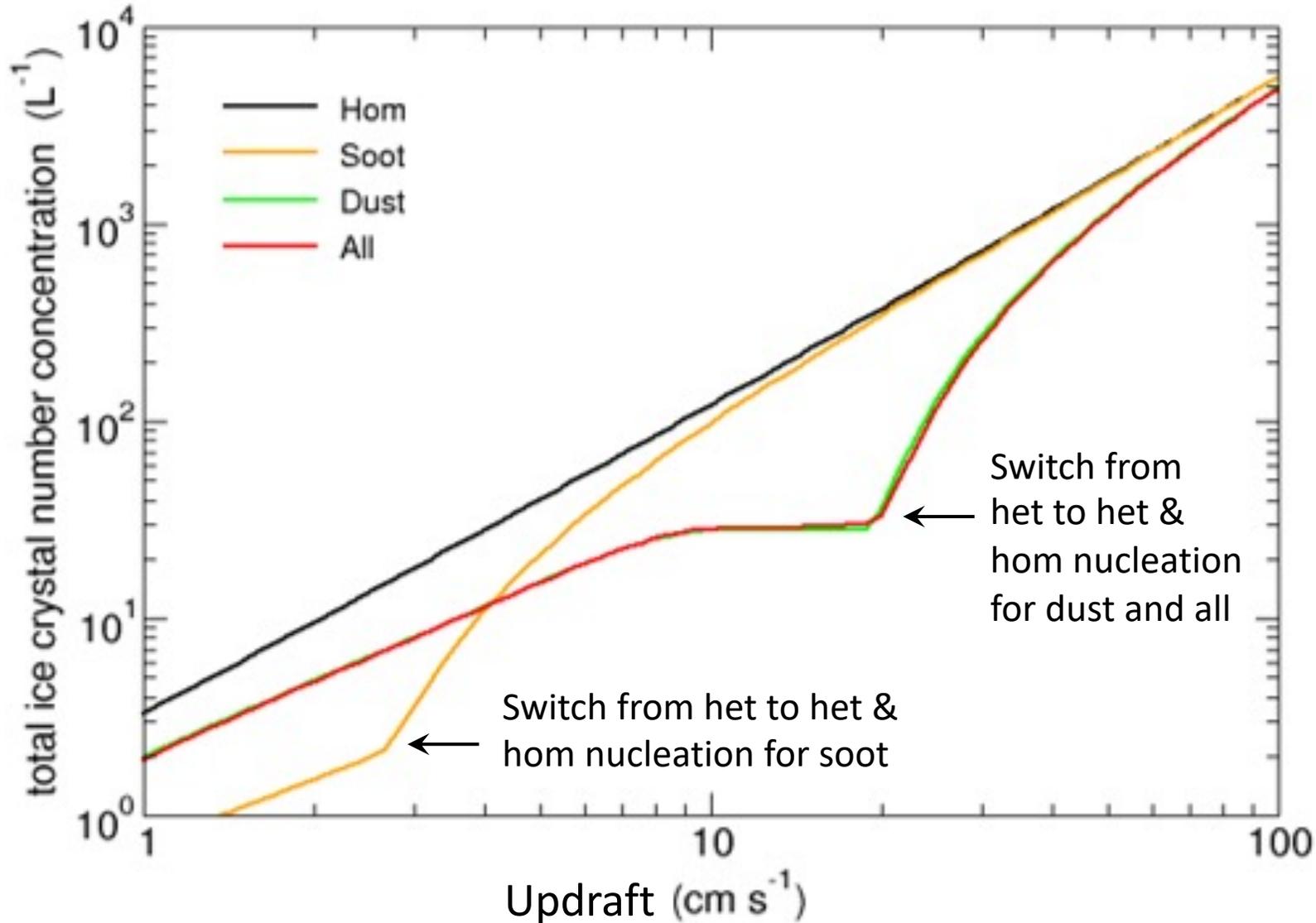
Ice-active soot particle size distribution



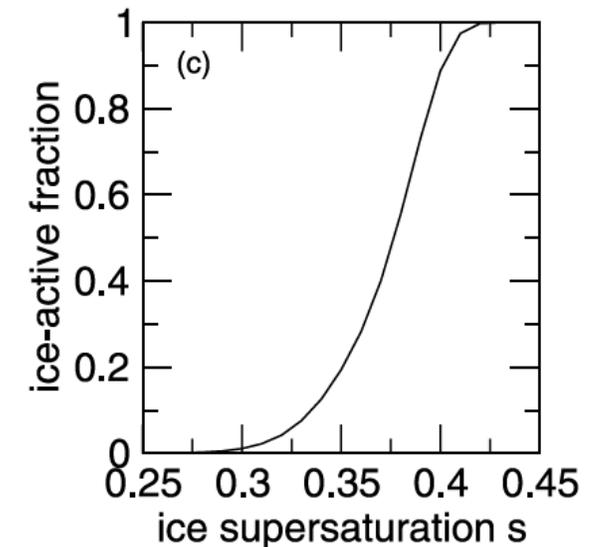
## Ice nucleation activity of soot at cirrus conditions assuming:

- bare soot aggregates (no coating)
- a distribution of primary particle sizes from 5–40 nm with mean  $D_{pp} = 20 \text{ nm}$
- Primary soot particle overlap from 0.01–0.2
- soot-water contact angle of  $60^\circ$

# Ice crystal formation at constant updraft



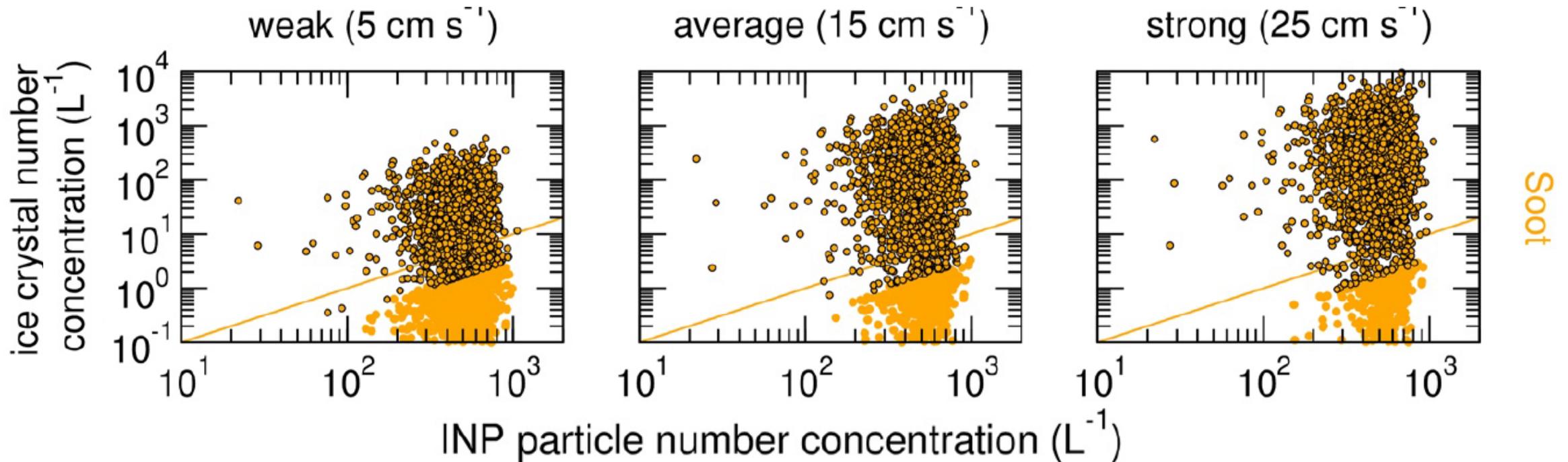
- 220 K, 250 hPa
- Liquid **background** aerosol particles:  $500\ 000\ L^{-1}$
- **Soot**:  $500\ L^{-1}$  but only about 1 % are able to activate
- **Background dust**:  $28\ L^{-1}$ , all activate at about  $s = 0.4$



# Ensemble simulations with high-resolution cirrus column model

- Typical upper tropospheric temperature (**220 K**) and pressure (**250 hPa**).
- Variable updraft speeds by random sampling from exponential distributions with standard deviations of **5 (weak), 15 (average), and 25 (strong) cm/s**.
- **Ice nucleation:**
  - **Homogeneously** on **constant** background aerosol: **500 000 L<sup>-1</sup>** liquid 500 nm-particles (wet diameter)
  - On **variable** soot particles sampled from normal distributions in concentration (with  $n_s = \mathbf{500\ L^{-1}}$  and  $\sigma = 150\ L^{-1}$ ) and size (with  $D_m = \mathbf{29.3\ nm}$  and  $\sigma = 1.72\ L^{-1}$ )
  - On a **variable** background mineral dust particle concentration sampled from a normal distribution in concentration (with  $n_s = \mathbf{28\ L^{-1}}$  and  $\sigma = 12\ L^{-1}$ ) and constant size distribution.
- **Ice growth:** constant deposition coefficients of  $\alpha = 0.3$  (dust) and 0.7 (soot).

# Competition between homogeneous ice nucleation and ice nucleation on soot

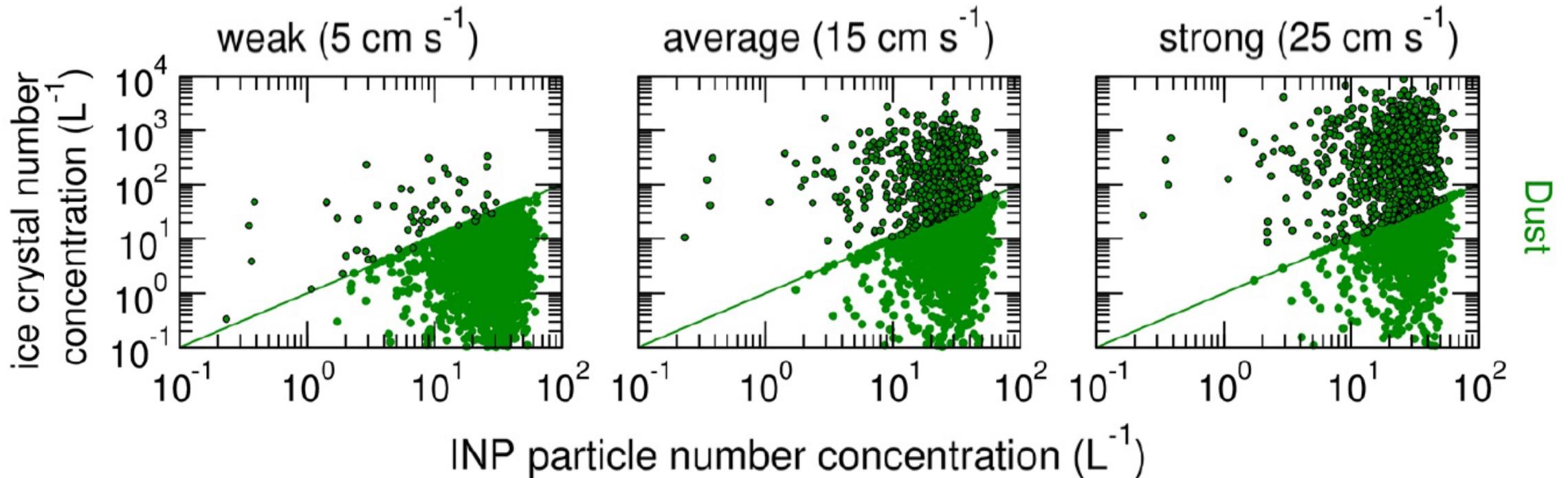


**Filled yellow circles:** ice formed only on soot particles (no homogeneous nucleation)

**Yellow circles with black outlines:** ice formed also homogeneously

**Yellow line:** ice crystal number concentration if all soot particles had activated to ice

# Competition between homogeneous ice nucleation and ice nucleation on mineral dust

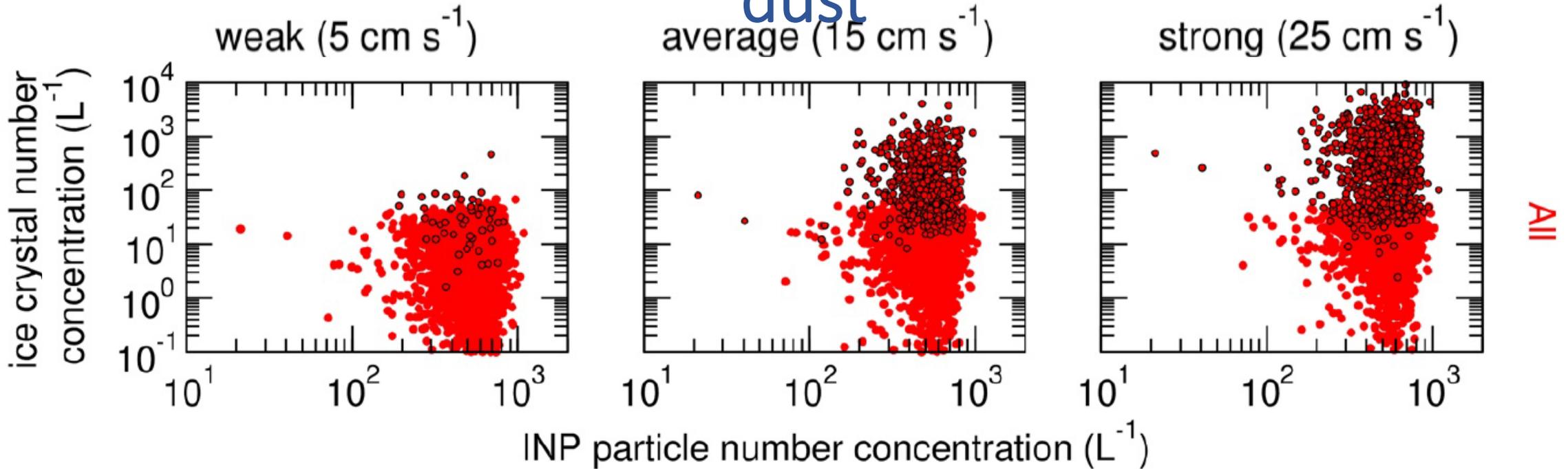


**Filled green circles:** ice formed only on background mineral dust (no homogeneous nucleation)

**Green circles with black outlines:** ice formed also homogeneously

**Green line:** ice crystal number concentration (ICNC) if all dust particles had activated to ice

# Competition between homogeneous ice nucleation and ice nucleation on soot and mineral dust



**Filled red circles:** ice formed only on soot or dust particles (no homogeneous nucleation)

**Red circles with black outlines:** ice formed also homogeneously

# Summary of all ensemble simulations

Case	Category	Updraft: 5 cm/s		Updraft: 15 cm/s		Updraft: 25 cm/s	
		Frequency	Median ICNC	Frequency	Median ICNC	Frequency	Median ICNC
Hom		100 %	14 L <sup>-1</sup>	100 %	74 L <sup>-1</sup>	100 %	164 L <sup>-1</sup>
Soot	only het	49 %	0.73 L <sup>-1</sup>	22 %	0.81 L <sup>-1</sup>	14 %	0.9 L <sup>-1</sup>
	Hom & het	51 %	22 L <sup>-1</sup>	78 %	92 L <sup>-1</sup>	86 %	215 L <sup>-1</sup>
Dust	only het	97 %	6 L <sup>-1</sup>	79 %	15 L <sup>-1</sup>	62 %	18 L <sup>-1</sup>
	Hom & het	3 %	21 L <sup>-1</sup>	21 %	151 L <sup>-1</sup>	38 %	272 L <sup>-1</sup>
All	only het	98 %	6 L <sup>-1</sup>	81 %	15 L <sup>-1</sup>	63 %	19 L <sup>-1</sup>
	Hom & het	2 %	22 L <sup>-1</sup>	19 %	164 L <sup>-1</sup>	37 %	275 L <sup>-1</sup>

→ soot does hardly influence ice crystal number concentration (ICNC) considering the omnipresence of a background dust concentration

# Conclusions

- Soot particles nucleate ice through a pore condensation and freezing process (PCF)
- Because of their tiny size, aircraft soot particles are poor INPs
- Upper tropospheric background concentrations of dust particles outcompete soot in cirrus cloud formation
- Aircraft soot does not seem to be relevant in cirrus cloud formation