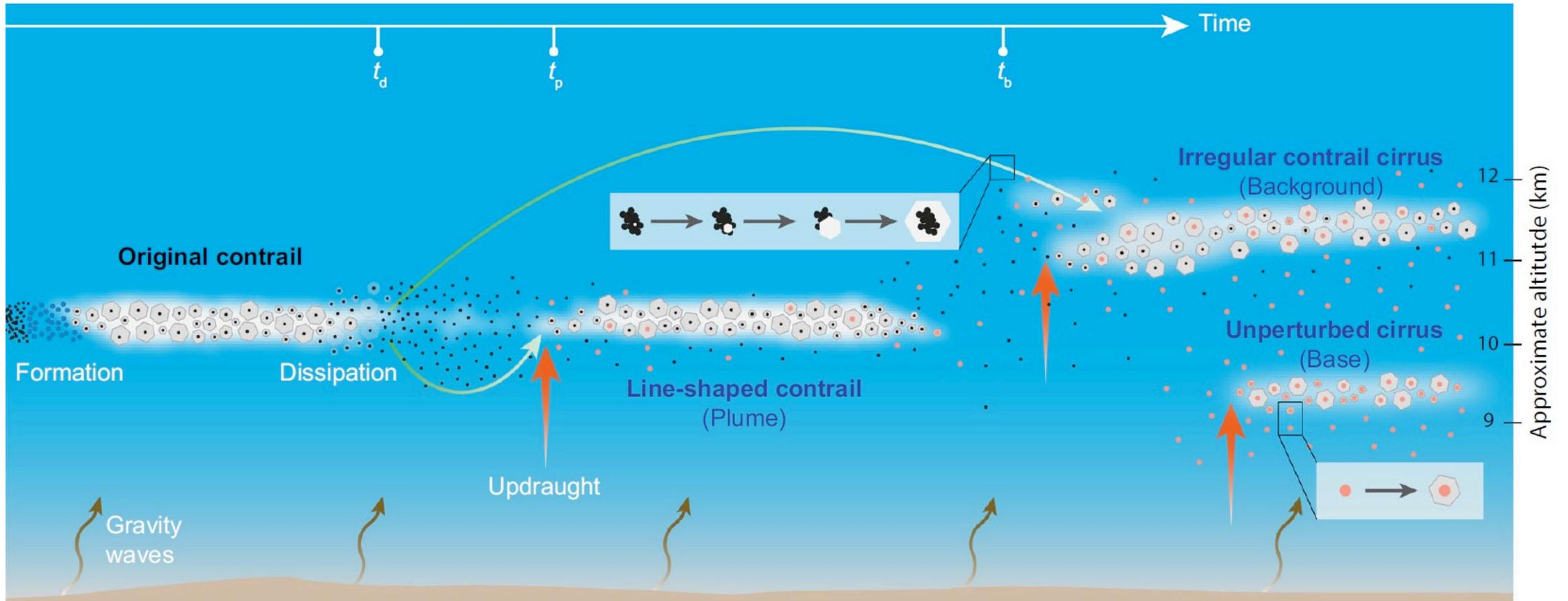


Heterogeneous ice nucleation of soot particles: measurements, predictions and implications

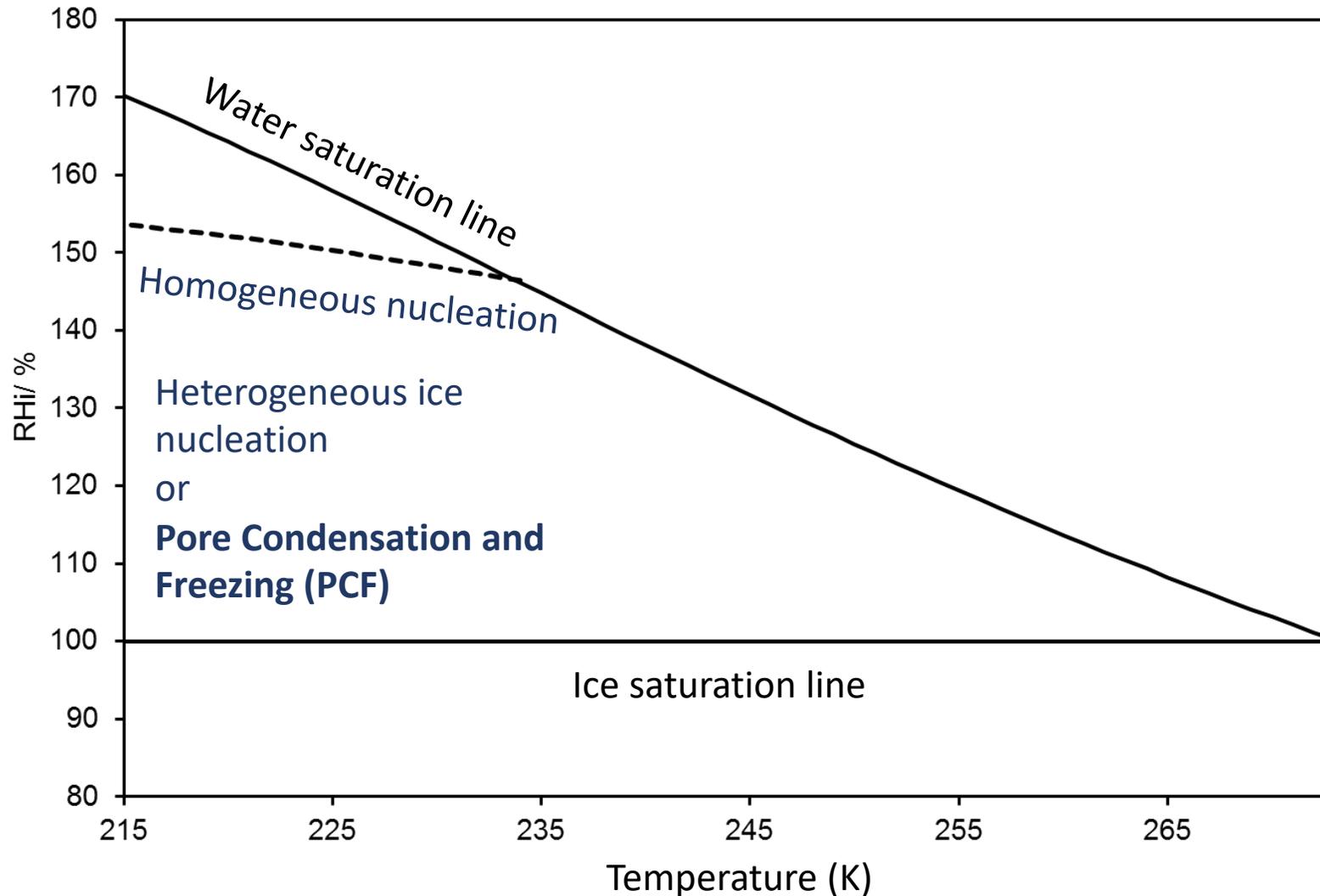
Claudia Marcolli¹, Fabian Mahrt², Bernd Kärcher³

¹ETH Zürich, ²Paul Scherrer Institute, ³DLR Oberpfaffenhofen

Overview



RH/T conditions for the different ice nucleation modes



Homogeneous ice nucleation:

Occurring on ubiquitous solution droplets

Heterogeneous ice nucleation:

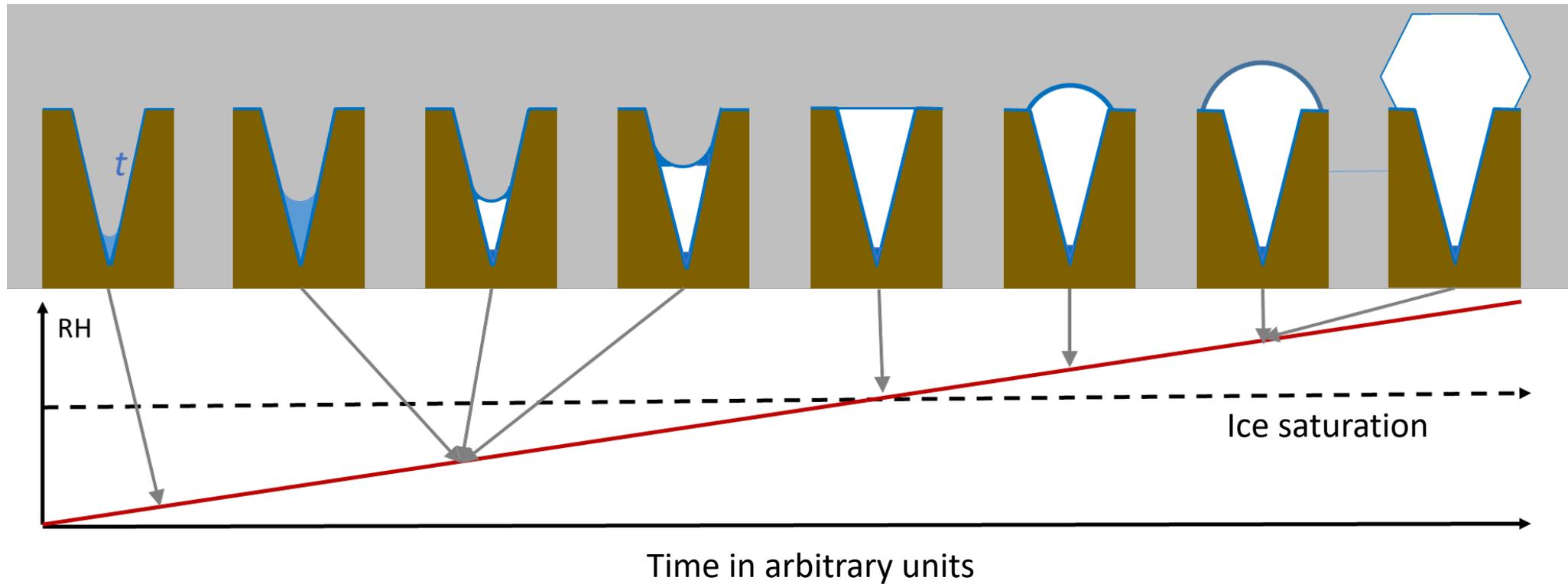
Occurring on ice-nucleating particles (INPs) with nucleation sites

PCF:

Occurring on porous particles through homogeneous ice nucleation of pore water

Pore condensation and freezing (PCF)

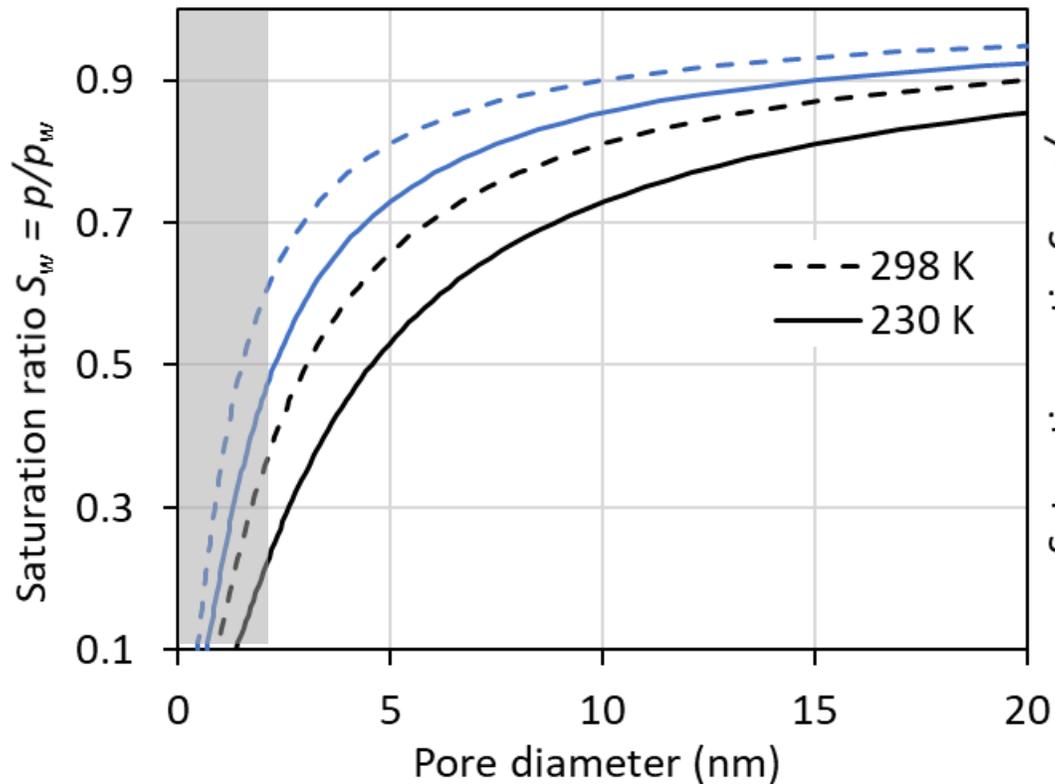
- Capillary condensation of water into pores (Kelvin equation)
- Nucleation of ice within the pores (Classical nucleation theory)
- Growth of ice out of the pores (Kelvin equation)



What is the relevant diameter range for PCF?

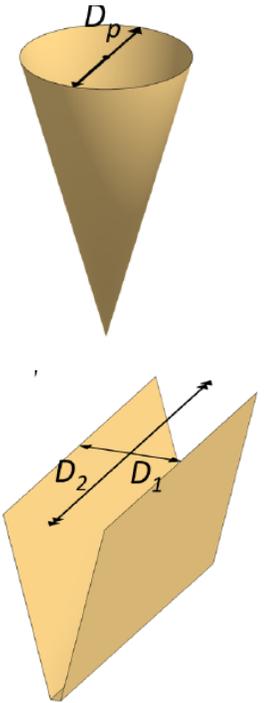
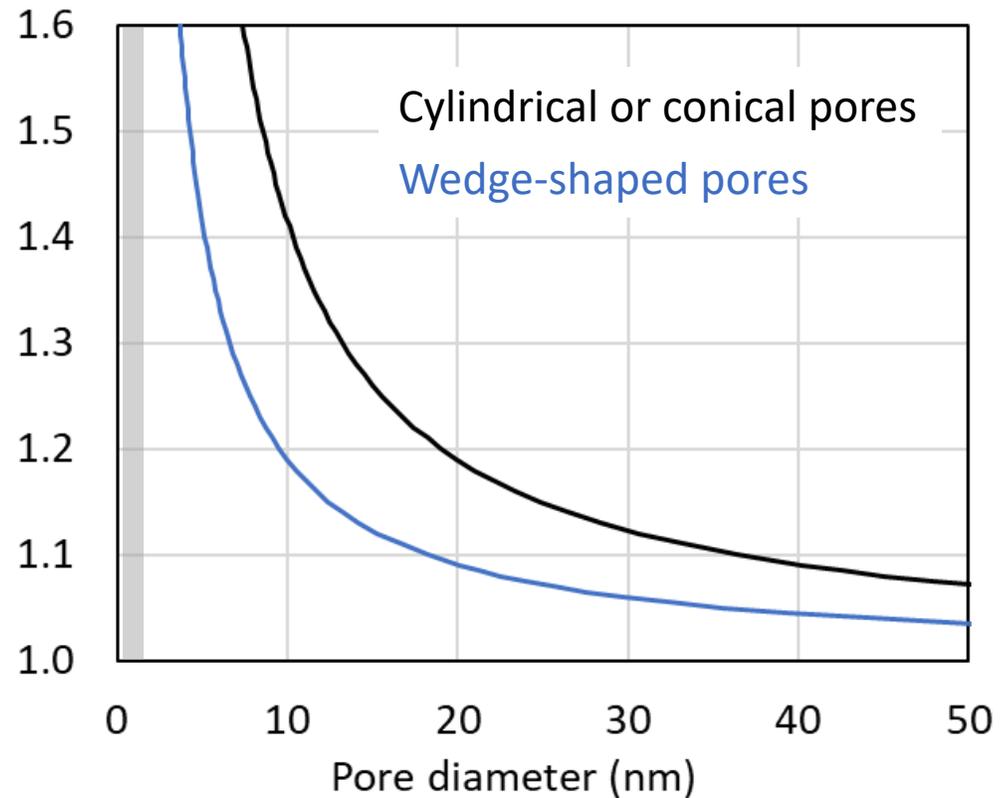
Pore filling with water

Too narrow to freeze

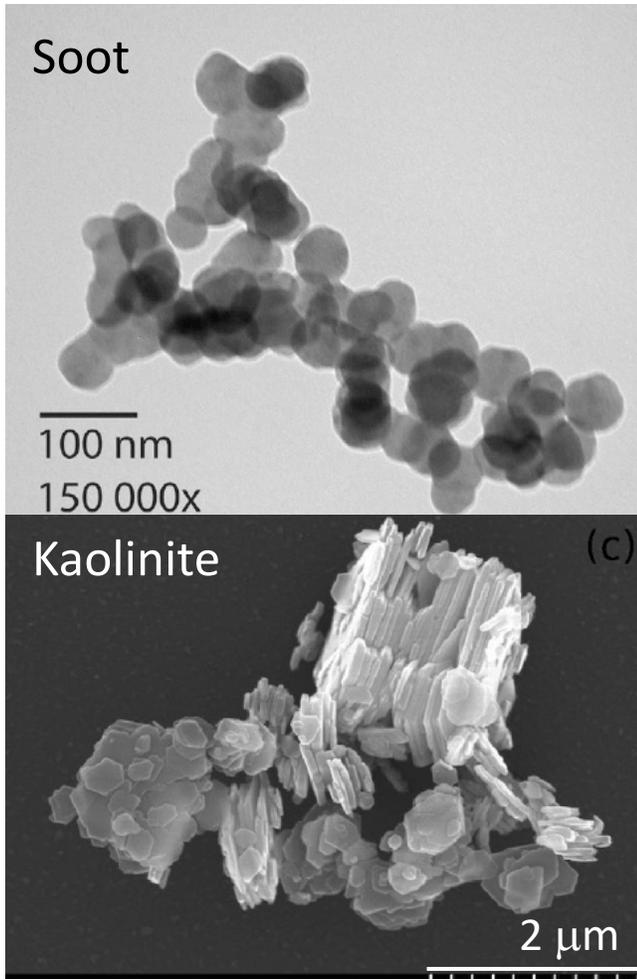


Ice growth out of the pores

Too narrow to freeze



What is the porosity of aerosol particles?



Soot

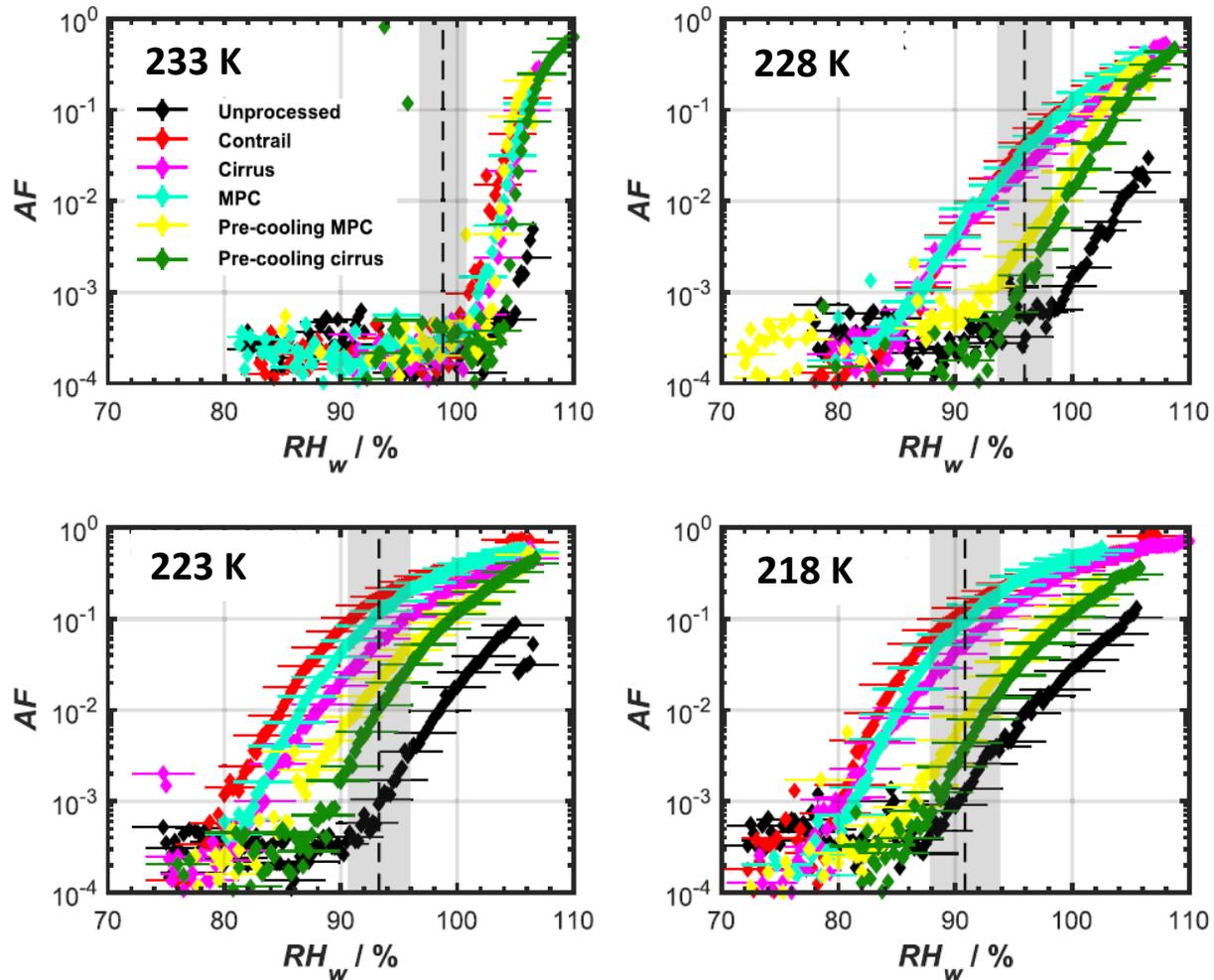
Porosity arising between primary particles

Mineral dust

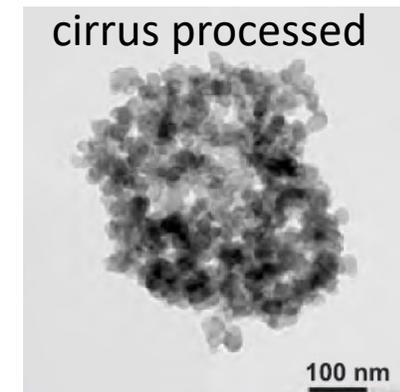
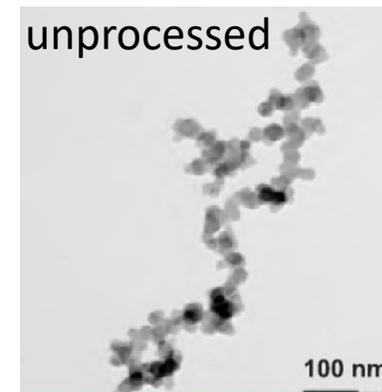
Porosity arising:

- in aggregates of crystallites
- At edges of clay minerals

Activated fraction of processed miniCAST black

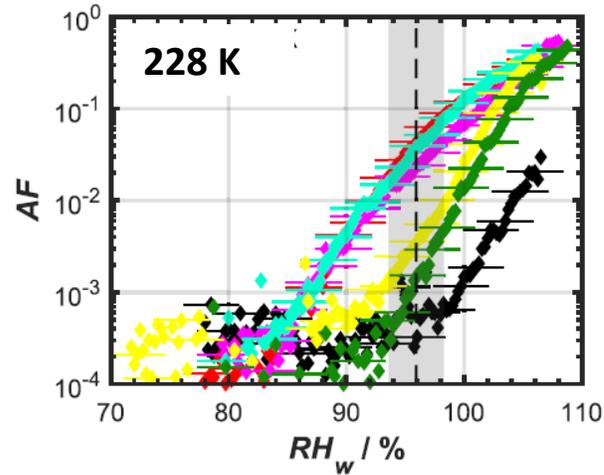
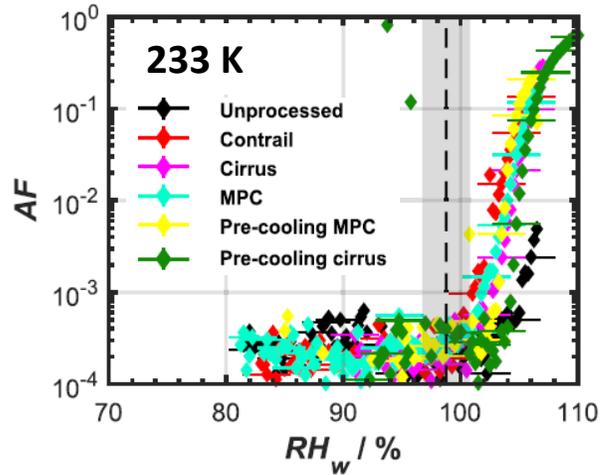


	Preprocessing	Evaporation
Unprocessed	—	—
contrail	104 % RH_w 228 K	38 % RH_w 233 K
cirrus	96 % RH_w 228 K	38 % RH_w 233 K
MPC	108 % RH_w 243 K	30 % RH_w 253 K
Precool MPC	96 % RH_w 243 K	30 % RH_w 253 K
Precool cirrus	65 % RH_w 228 K	38 % RH_w 233 K

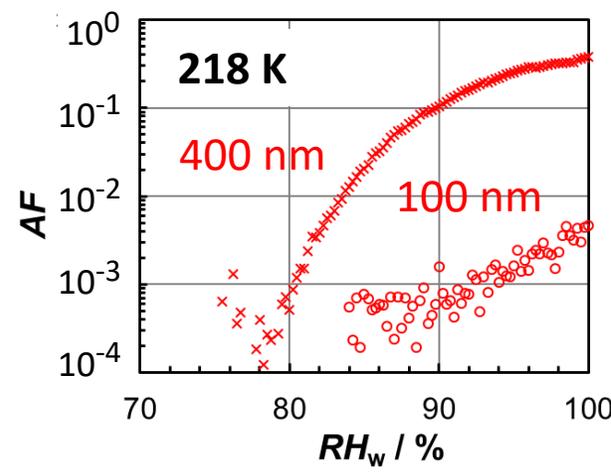
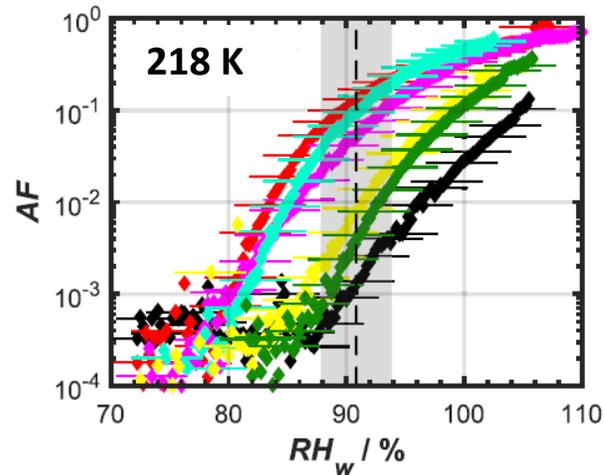
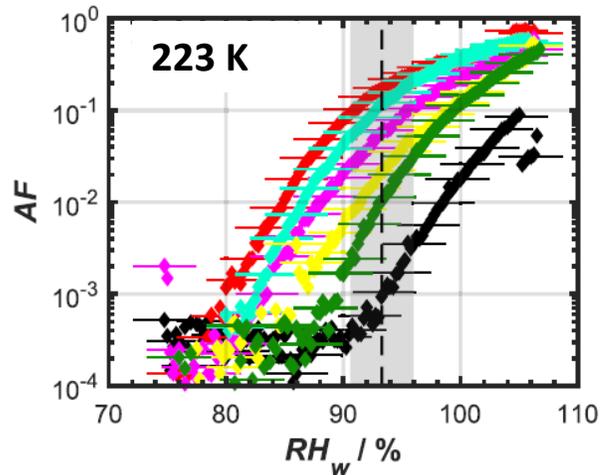
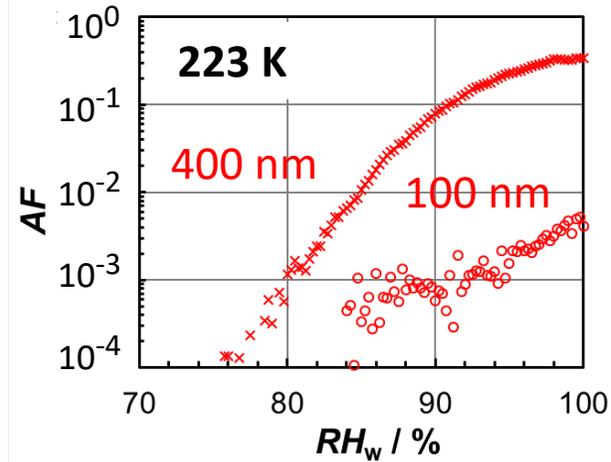


Activated fraction of processed miniCAST black

400 nm diameter

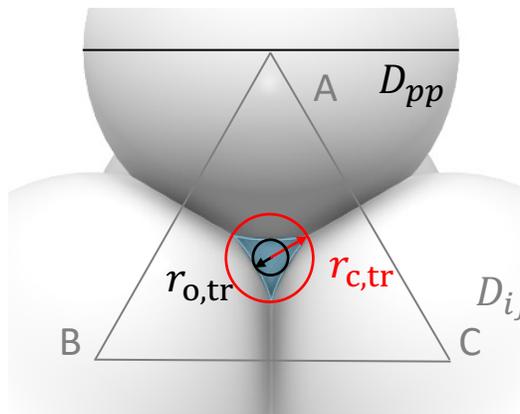
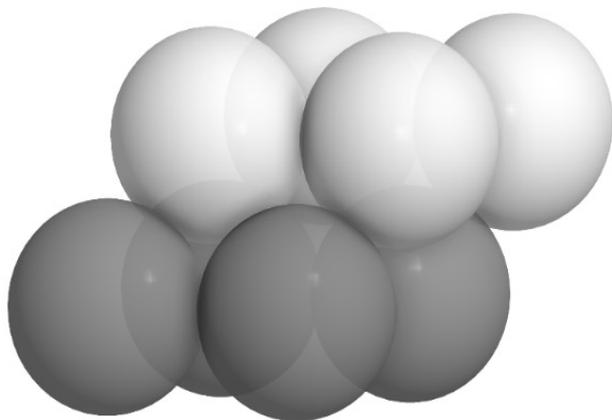


Size dependence



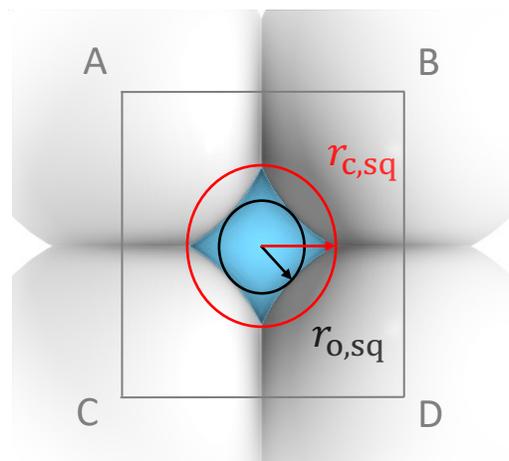
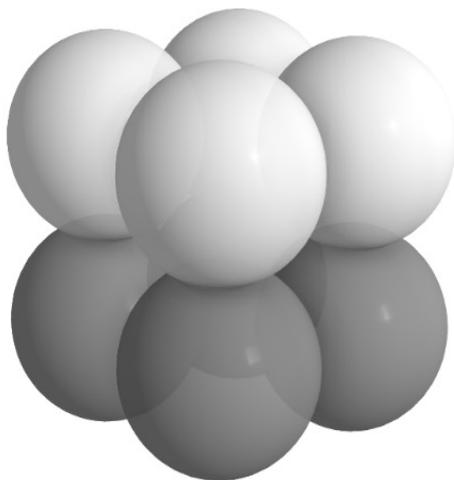
Pores within soot aggregates

Tetrahedral packing



Three-membered ring pore

Cubic packing



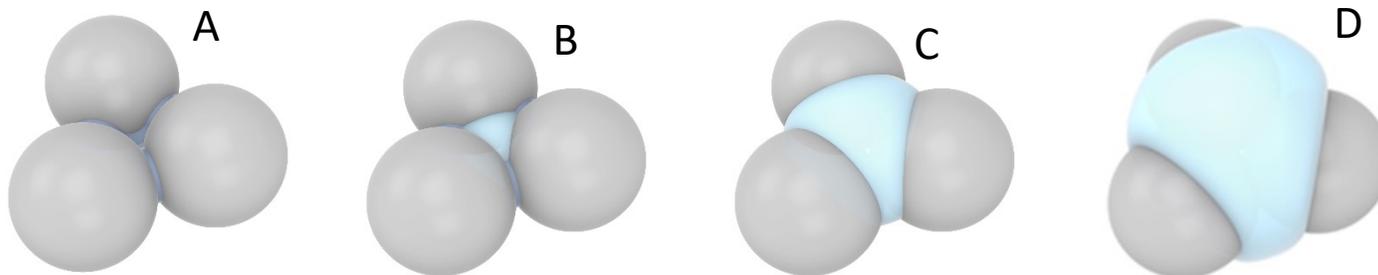
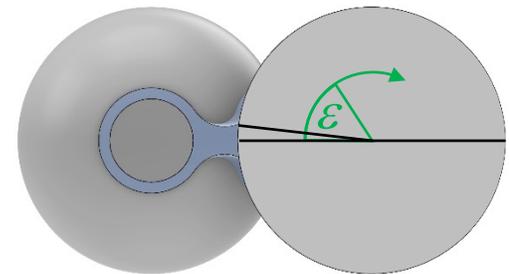
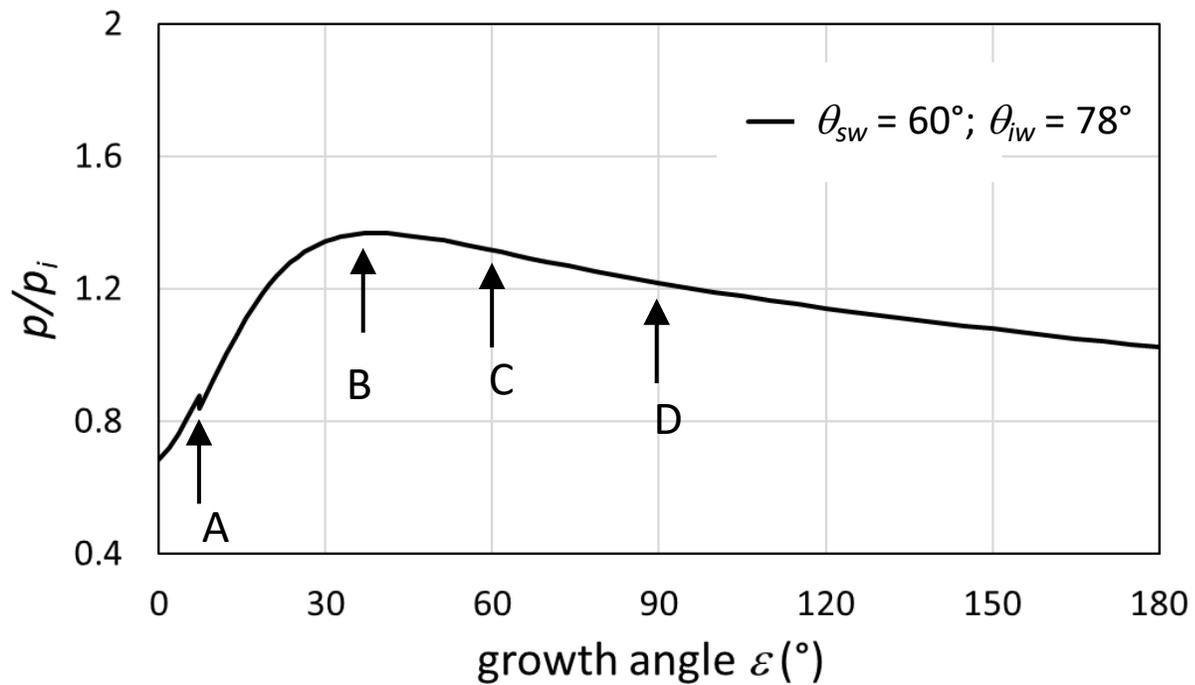
Four-membered ring pore

Spherical primary particles with typical diameters $D_{pp} = 10\text{--}30$ nm

$$\text{Overlap } C_{ov} = \frac{D_{pp} - D_{ij}}{D_{pp}} = 0.01\text{--}0.2$$

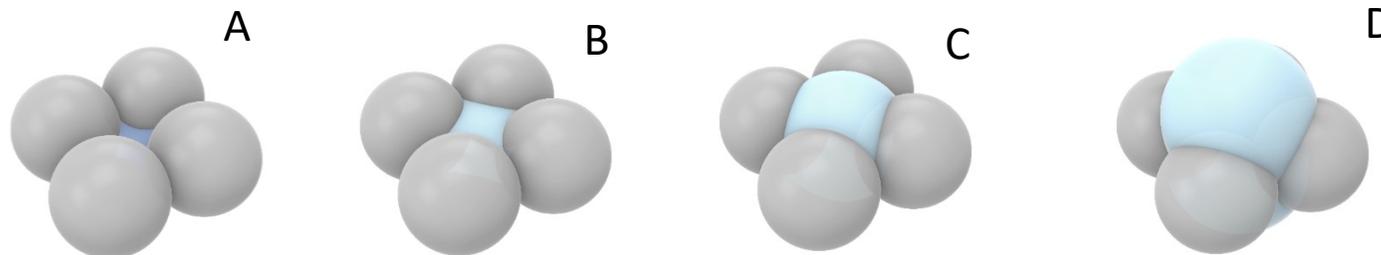
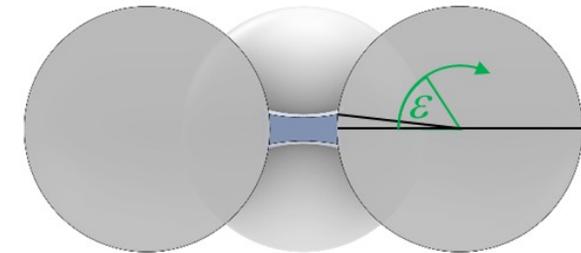
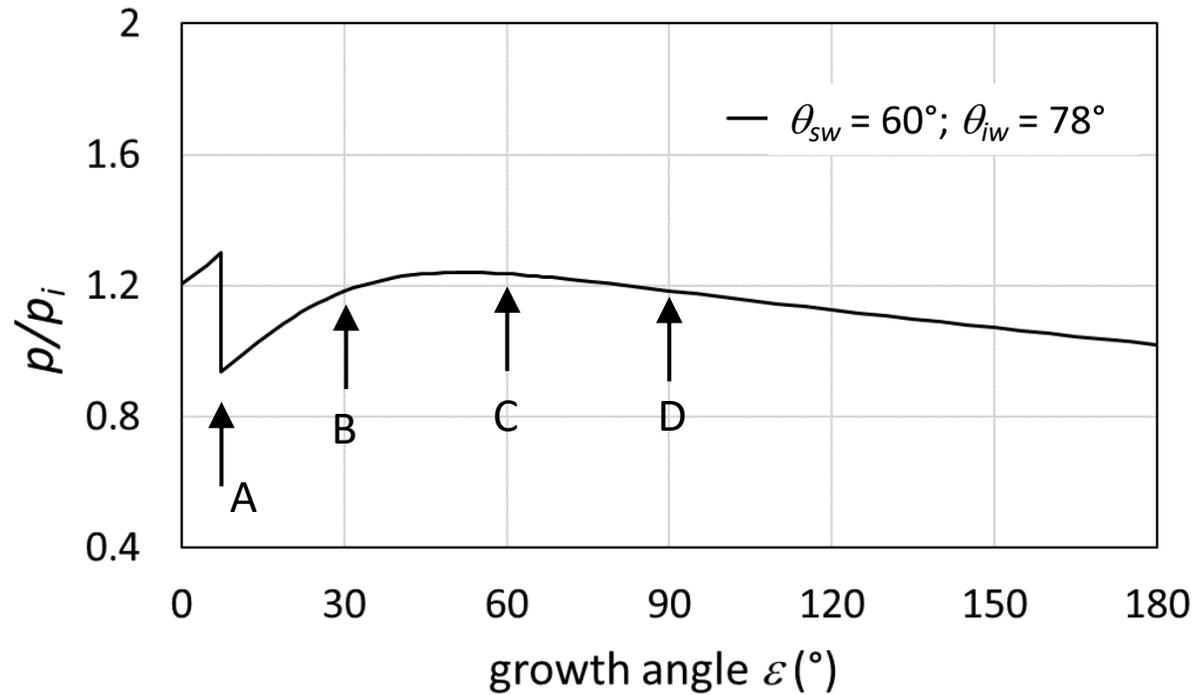
Soot PCF on three-membered ring pore

primary particle diameter: 20 nm, $C_{ov} = 0.05$

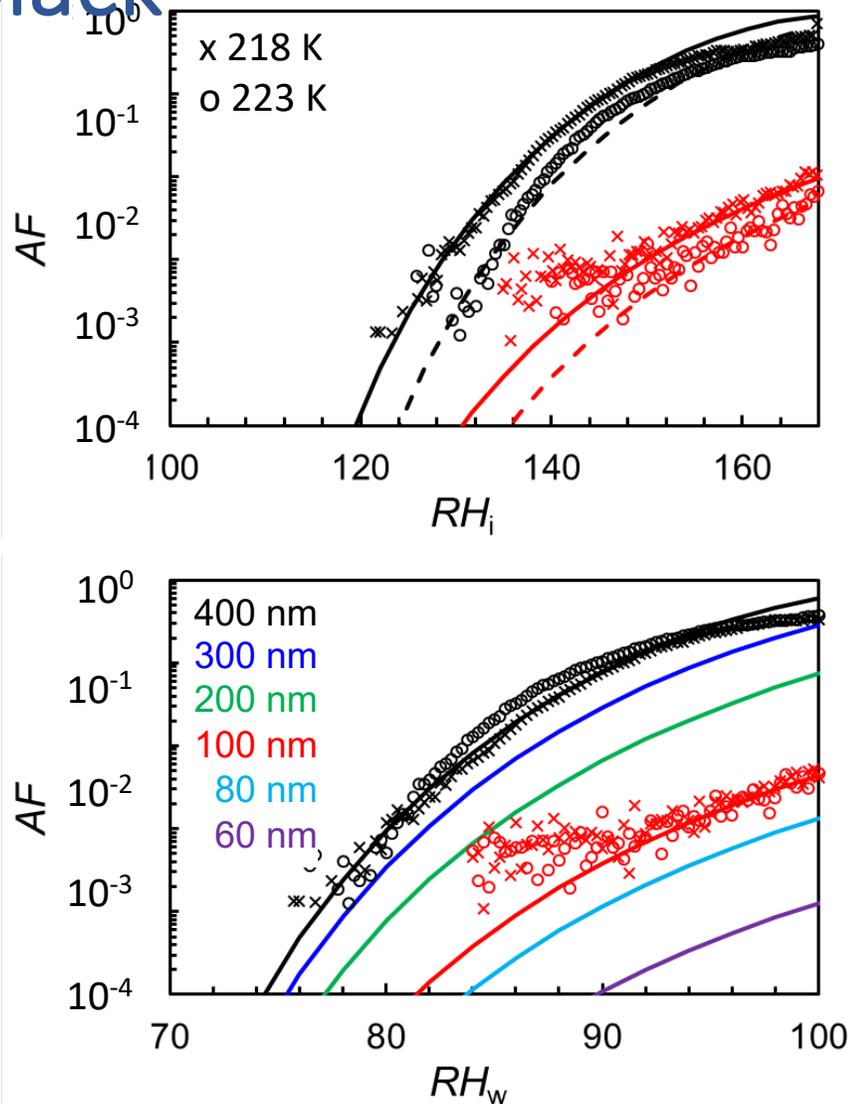


Soot PCF on four-membered ring pore

primary particle diameter: 20 nm, $C_{ov} = 0.1$



Activated fraction (AF) parameterization of miniCAST black



$$AF = 1 - (1 - P_N(RH)) \left((N_p - 2)^{1.86} \right)$$

$$P_N(RH_w) = 10^{\left(\frac{1}{0.3374 - 0.006091 RH_w} \right)}$$

Primary particle diameter: $D_{pp} = 31 \text{ nm}$

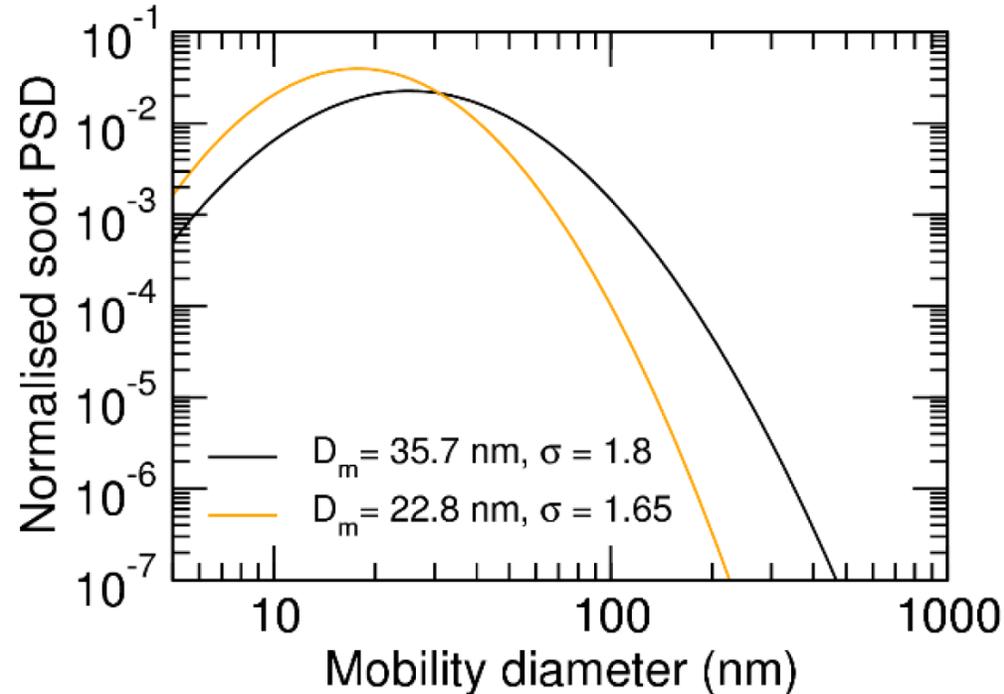
Fractal dimension: 1.86

→ Number of particles (N_p) in aggregate with diameter D_m

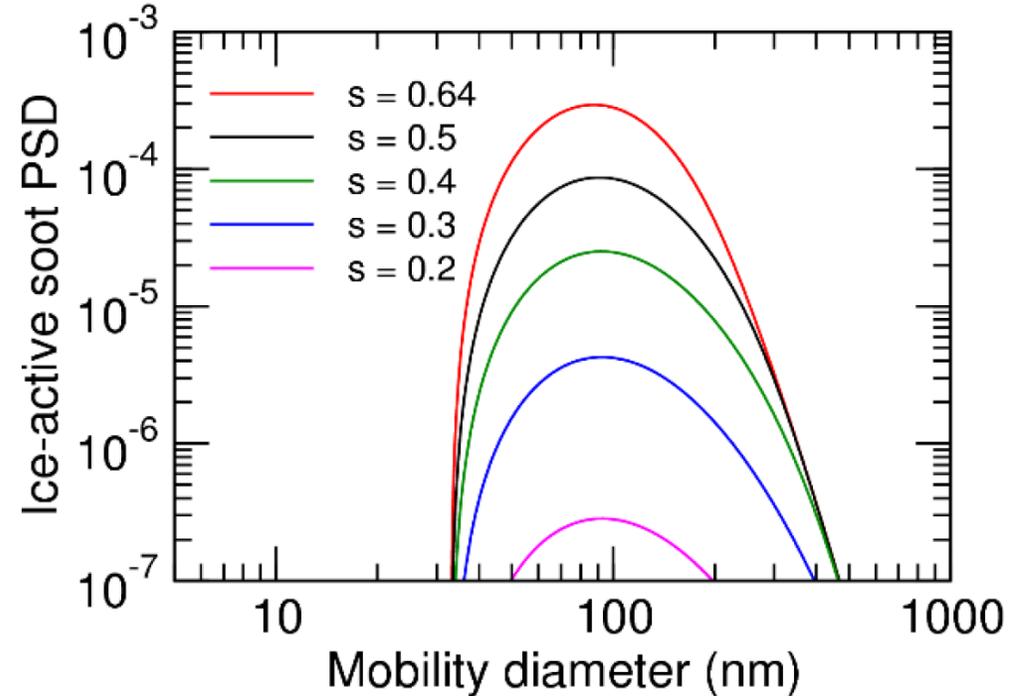
D_m	N_p
400 nm	94
300 nm	55
200 nm	26
100 nm	7
80 nm	5
60 nm	3

PCF parameterization for aircraft soot

Aircraft soot particle size distribution



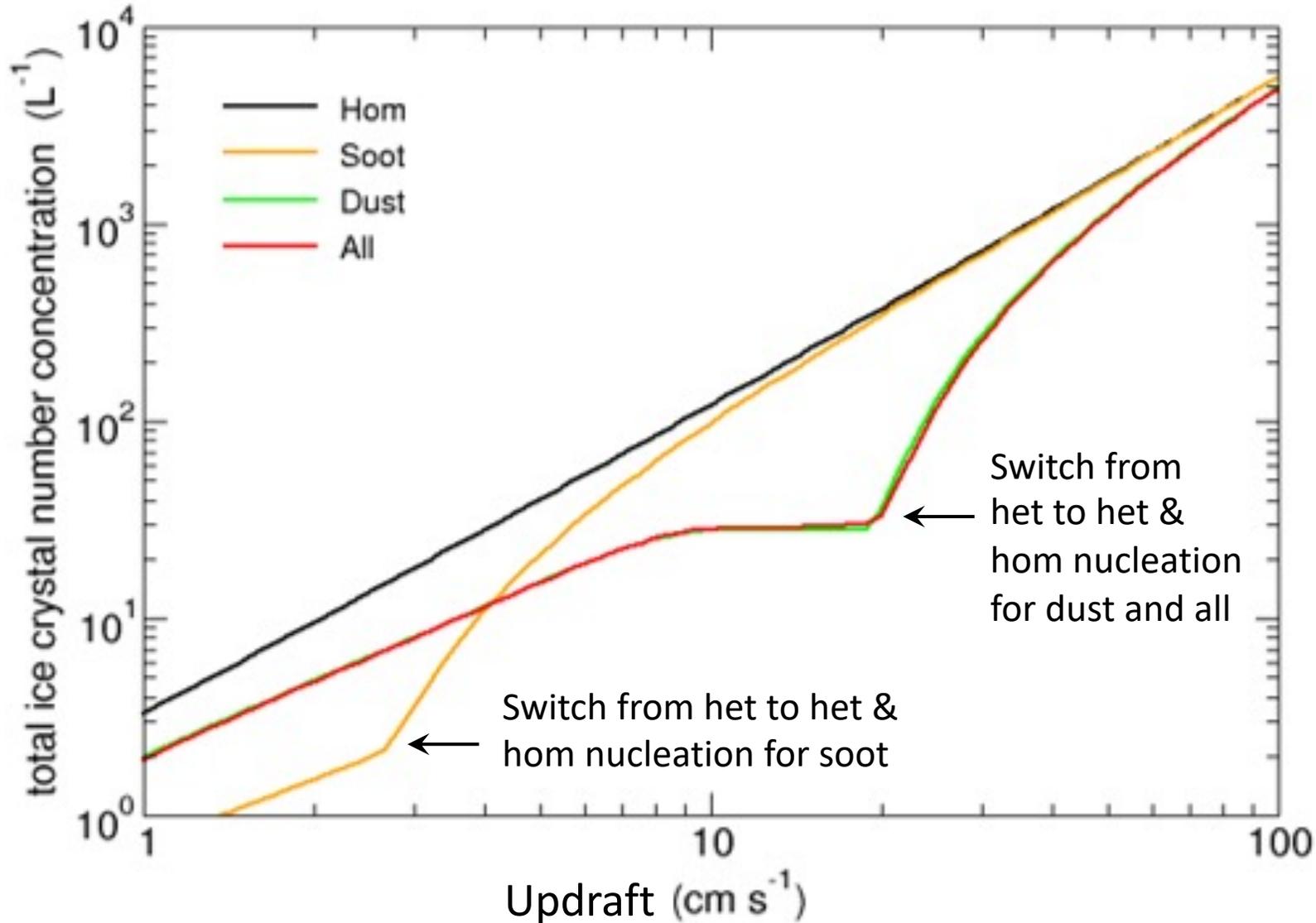
Ice-active soot particle size distribution



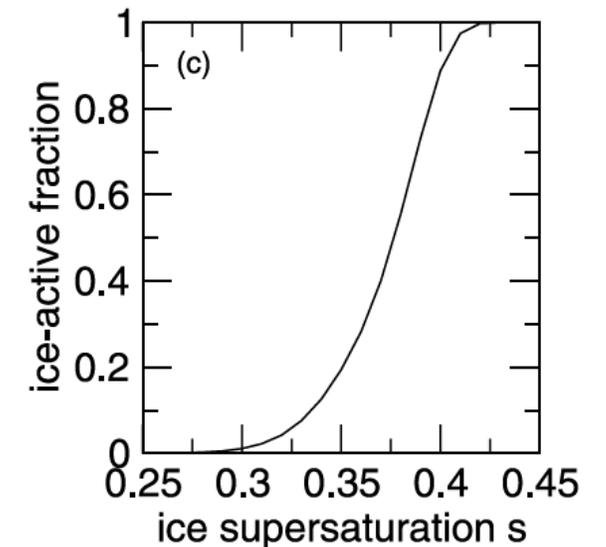
Ice nucleation activity of soot at cirrus conditions assuming:

- bare soot aggregates (no coating)
- a distribution of primary particle sizes from 5–40 nm with mean $D_{pp} = 20 \text{ nm}$
- Primary soot particle overlap from 0.01–0.2
- soot-water contact angle of 60°

Ice crystal formation at constant updraft



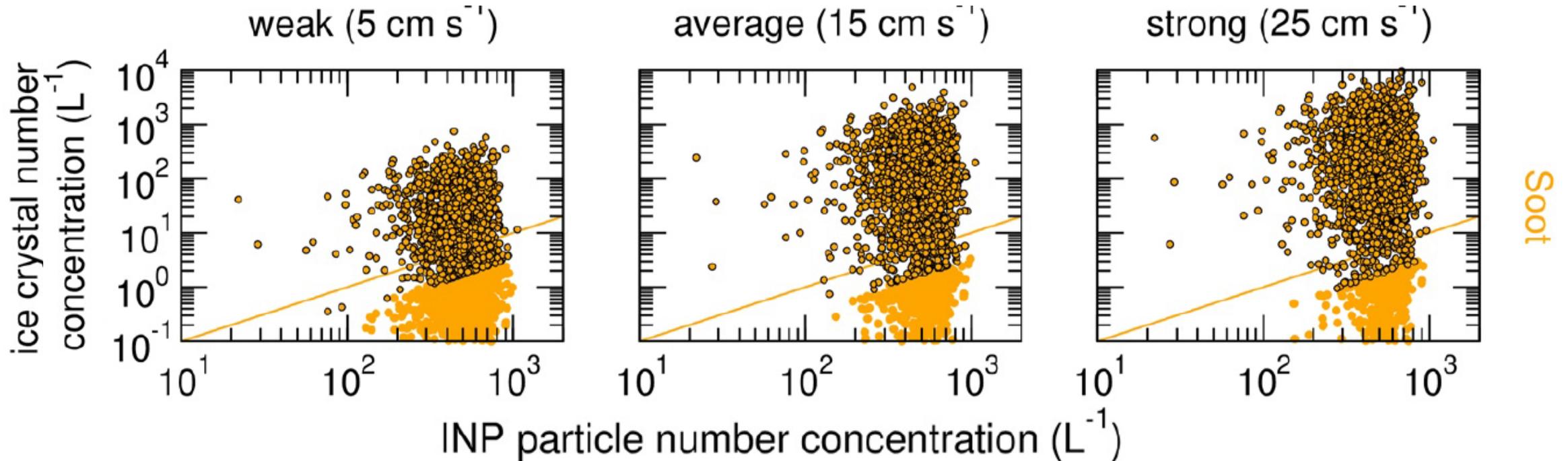
- 220 K, 250 hPa
- Liquid **background** aerosol particles: $500\ 000\ L^{-1}$
- **Soot**: $500\ L^{-1}$ but only about 1 % are able to activate
- **Background dust**: $28\ L^{-1}$, all activate at about $s = 0.4$



Ensemble simulations with high-resolution cirrus column model

- Typical upper tropospheric temperature (**220 K**) and pressure (**250 hPa**).
- Variable updraft speeds by random sampling from exponential distributions with standard deviations of **5 (weak)**, **15 (average)**, and **25 (strong) cm/s**.
- **Ice nucleation:**
 - **Homogeneously** on **constant** background aerosol: **500 000 L⁻¹** liquid 500 nm-particles (wet diameter)
 - On **variable** soot particles sampled from normal distributions in concentration (with $n_s = \mathbf{500\ L^{-1}}$ and $\sigma = 150\ L^{-1}$) and size (with $D_m = \mathbf{29.3\ nm}$ and $\sigma = 1.72\ L^{-1}$)
 - On a **variable** background mineral dust particle concentration sampled from a normal distribution in concentration (with $n_s = \mathbf{28\ L^{-1}}$ and $\sigma = 12\ L^{-1}$) and constant size distribution.
- **Ice growth:** constant deposition coefficients of $\alpha = 0.3$ (dust) and 0.7 (soot).

Competition between homogeneous ice nucleation and ice nucleation on soot

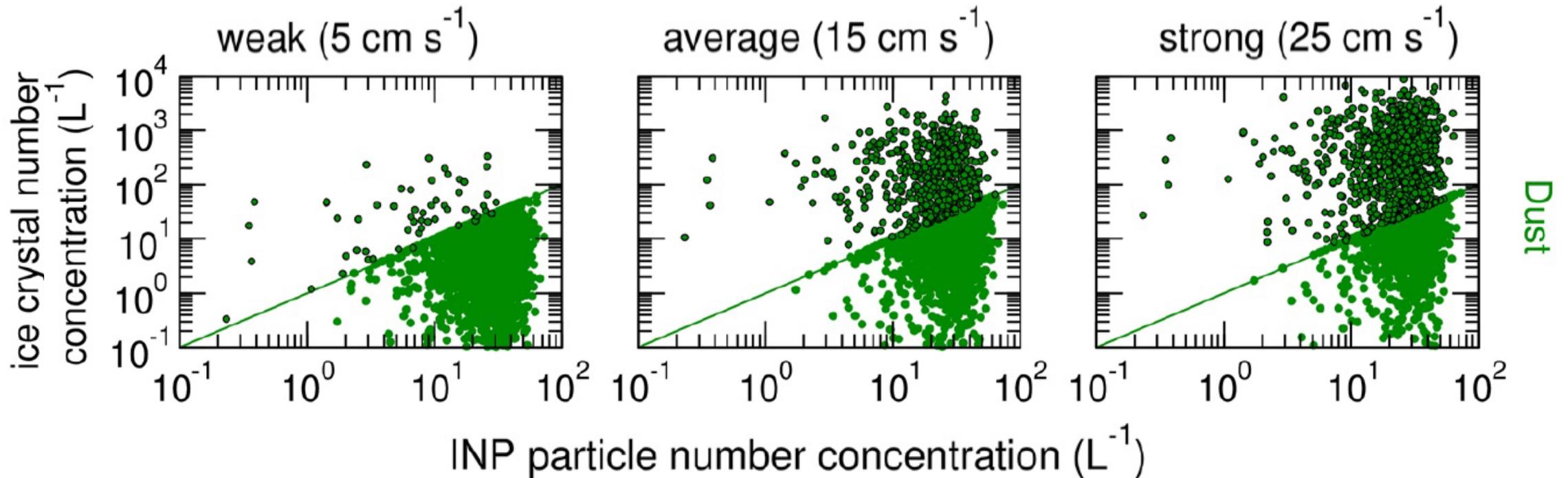


Filled yellow circles: ice formed only on soot particles (no homogeneous nucleation)

Yellow circles with black outlines: ice formed also homogeneously

Yellow line: ice crystal number concentration if all soot particles had activated to ice

Competition between homogeneous ice nucleation and ice nucleation on mineral dust

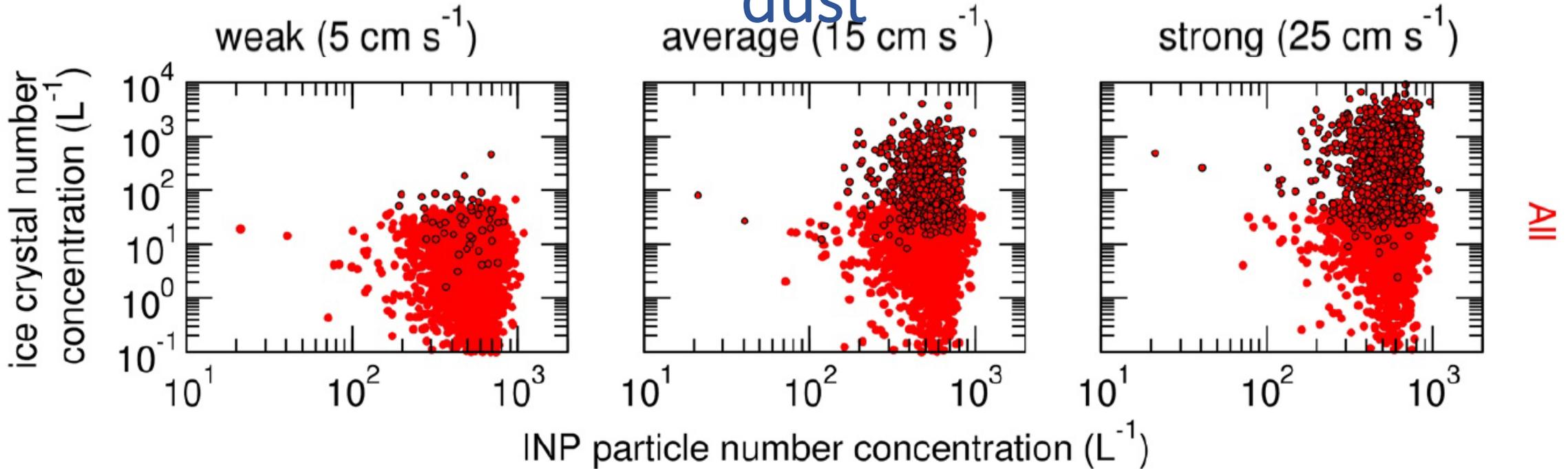


Filled green circles: ice formed only on background mineral dust (no homogeneous nucleation)

Green circles with black outlines: ice formed also homogeneously

Green line: ice crystal number concentration (ICNC) if all dust particles had activated to ice

Competition between homogeneous ice nucleation and ice nucleation on soot and mineral dust



Filled red circles: ice formed only on soot or dust particles (no homogeneous nucleation)

Red circles with black outlines: ice formed also homogeneously

Summary of all ensemble simulations

Case	Category	Updraft: 5 cm/s		Updraft: 15 cm/s		Updraft: 25 cm/s	
		Frequency	Median ICNC	Frequency	Median ICNC	Frequency	Median ICNC
Hom		100 %	14 L ⁻¹	100 %	74 L ⁻¹	100 %	164 L ⁻¹
Soot	only het	49 %	0.73 L ⁻¹	22 %	0.81 L ⁻¹	14 %	0.9 L ⁻¹
	Hom & het	51 %	22 L ⁻¹	78 %	92 L ⁻¹	86 %	215 L ⁻¹
Dust	only het	97 %	6 L ⁻¹	79 %	15 L ⁻¹	62 %	18 L ⁻¹
	Hom & het	3 %	21 L ⁻¹	21 %	151 L ⁻¹	38 %	272 L ⁻¹
All	only het	98 %	6 L ⁻¹	81 %	15 L ⁻¹	63 %	19 L ⁻¹
	Hom & het	2 %	22 L ⁻¹	19 %	164 L ⁻¹	37 %	275 L ⁻¹

→ soot does hardly influence ice crystal number concentration (ICNC) considering the omnipresence of a background dust concentration

Conclusions

- Soot particles nucleate ice through a pore condensation and freezing process (PCF)
- Because of their tiny size, aircraft soot particles are poor INPs
- Upper tropospheric background concentrations of dust particles outcompete soot in cirrus cloud formation
- Aircraft soot does not seem to be relevant in cirrus cloud formation